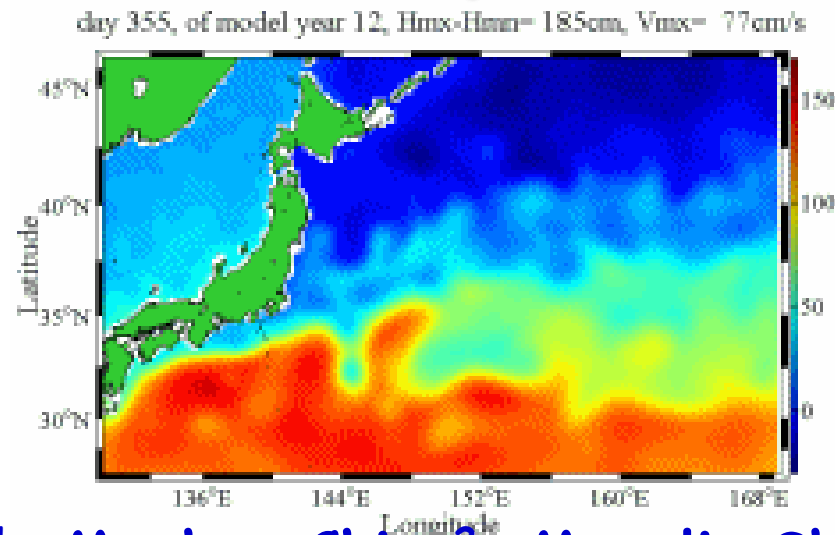


# Parallel domain-decomposed Taiwan Multi-scale Community Ocean Model (PD-TIMCOM)



Yu-Heng Tseng<sup>1</sup>, Mu-hua Chien<sup>2</sup>, Mao-lin Shen<sup>2</sup>, Chih-Chieh Young<sup>2</sup>

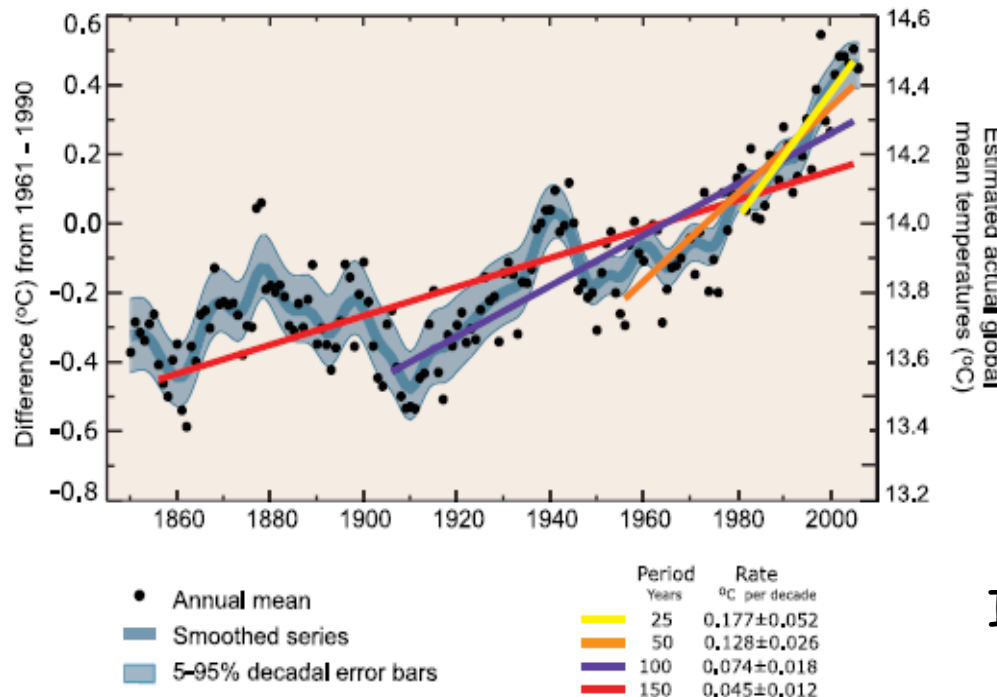
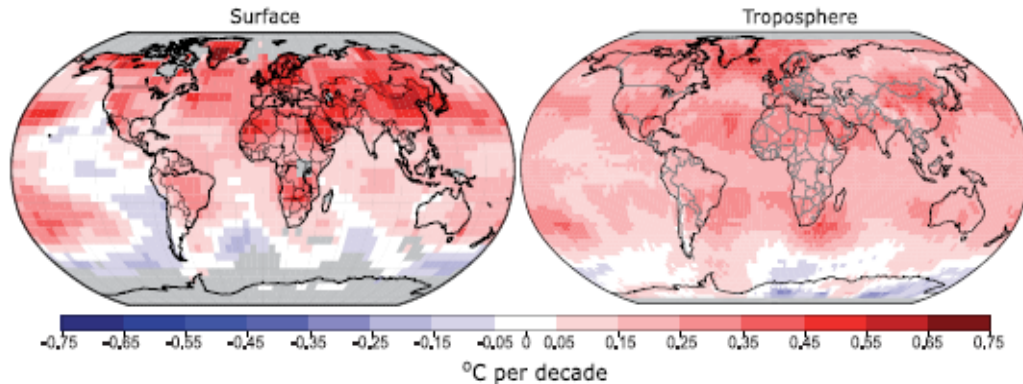
<sup>1</sup>Climate & Global Dynamics Division  
National Center for Atmospheric Research

<sup>2</sup>Department of Atmospheric Sciences  
National Taiwan University

Acknowledgement: computing resources from NCHC, Taiwan and NERSC, USA

# Climate change and ocean

## GLOBAL TEMPERATURE TRENDS



## GLACIAL-INTERGLACIAL ICE CORE DATA

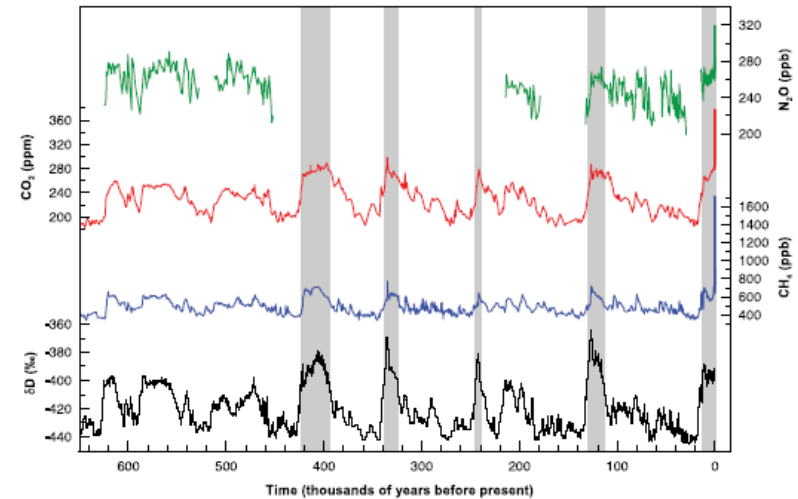
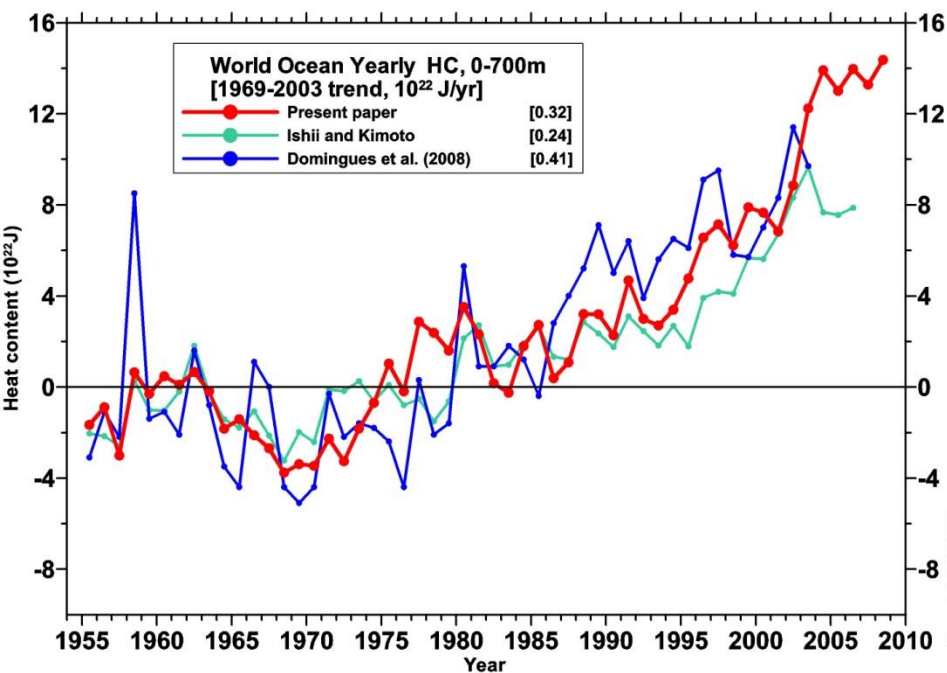
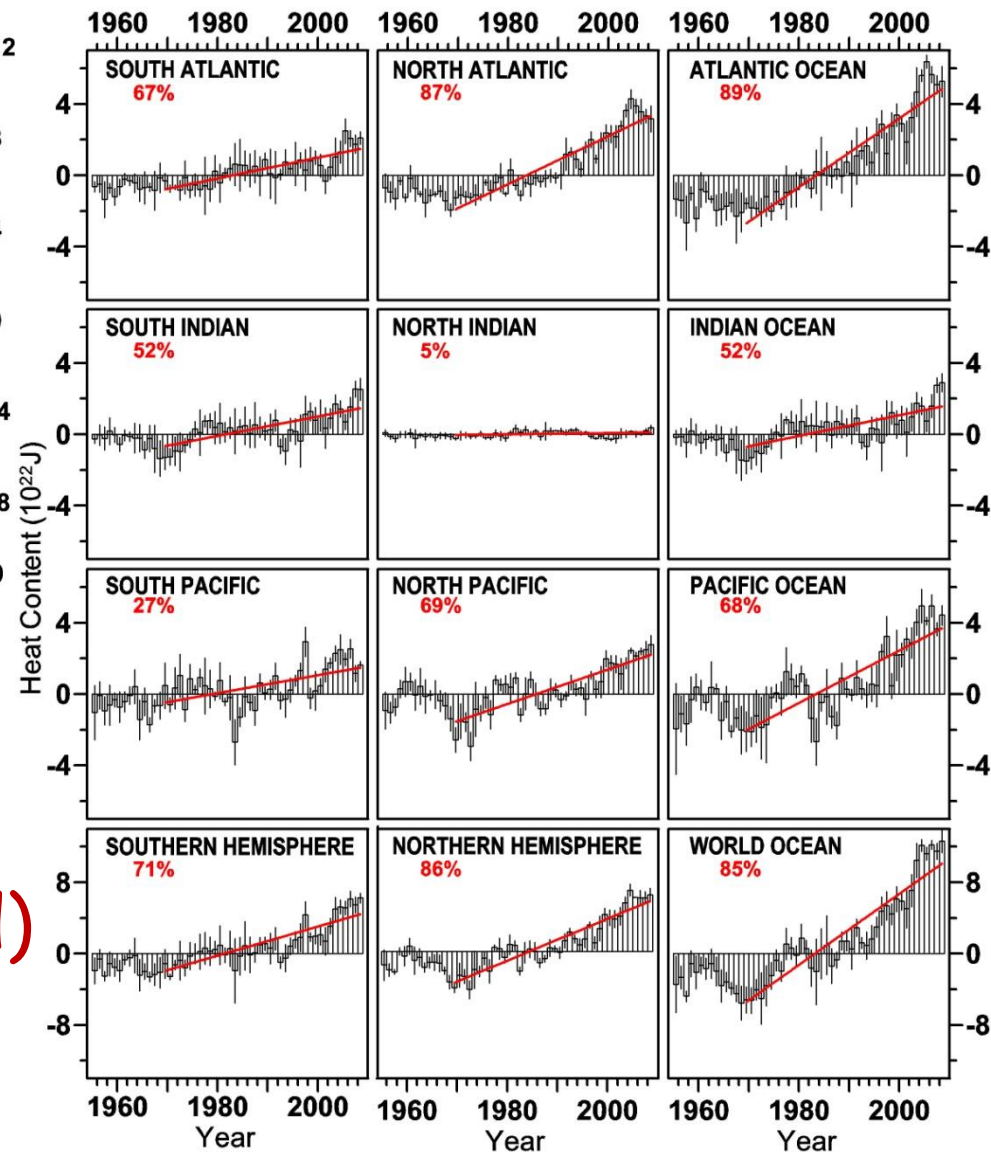


Figure TS.1. Variations of deuterium (δD) in antarctic ice, which is a proxy for local temperature, and the atmospheric concentrations the greenhouse gases carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>), and nitrous oxide (N<sub>2</sub>O) in air trapped within the ice core and from recent atmospheric measurements. Data cover 650,000 years and the shaded bands indicate current and previous interglacial warm period (Adapted from Figure 6.3)

IPCC AR4 WG1 Technical Summary



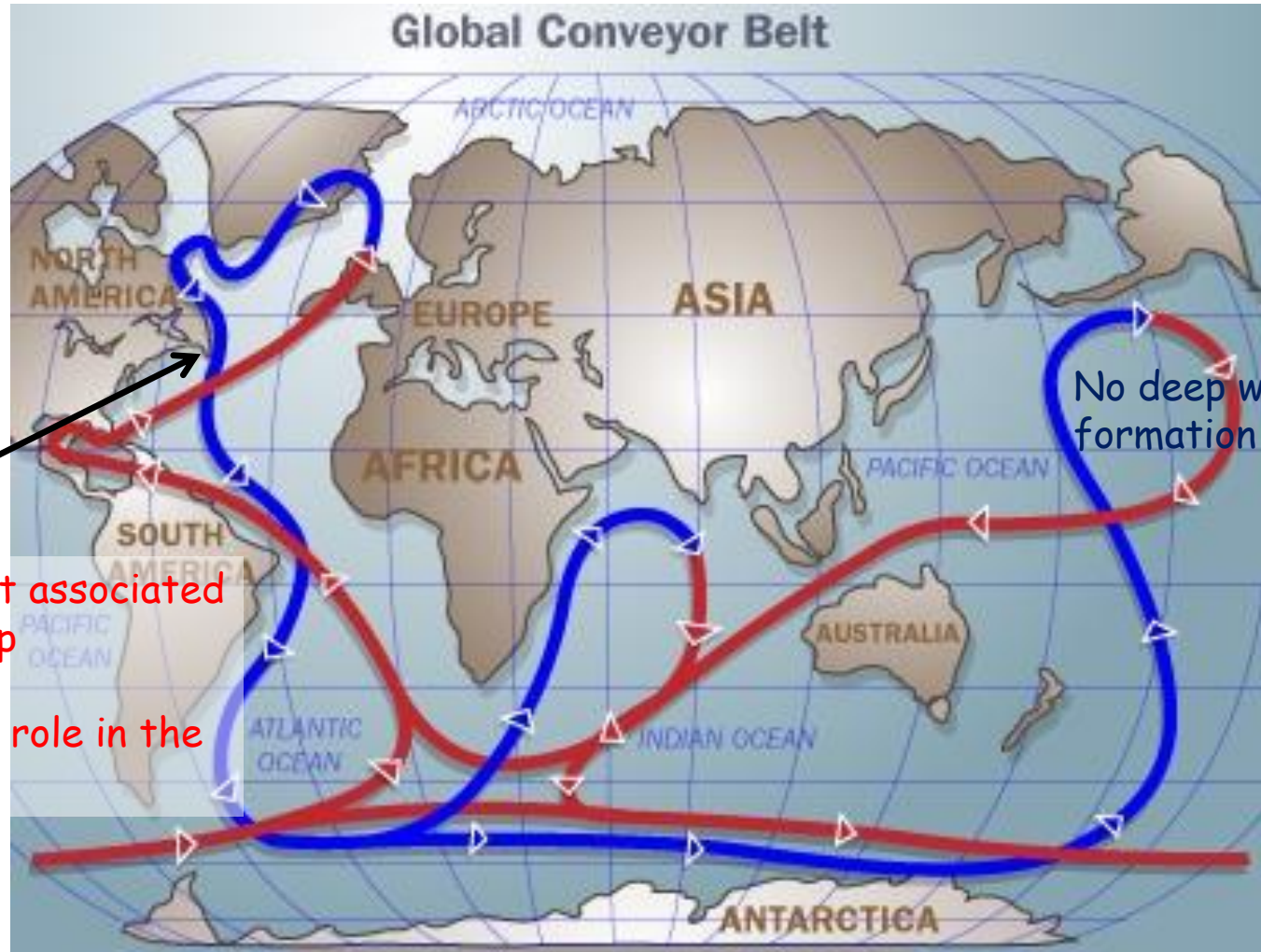
Global ocean variation  
(long term warming trend)



Levitus et al. (2009)

# Global Thermohaline Circulation:

- 1, thermal forcing, when water is cooled and sinks, and
- 2, haline forcing, when excess precipitation makes water less dense, and thus resistant to sinking.



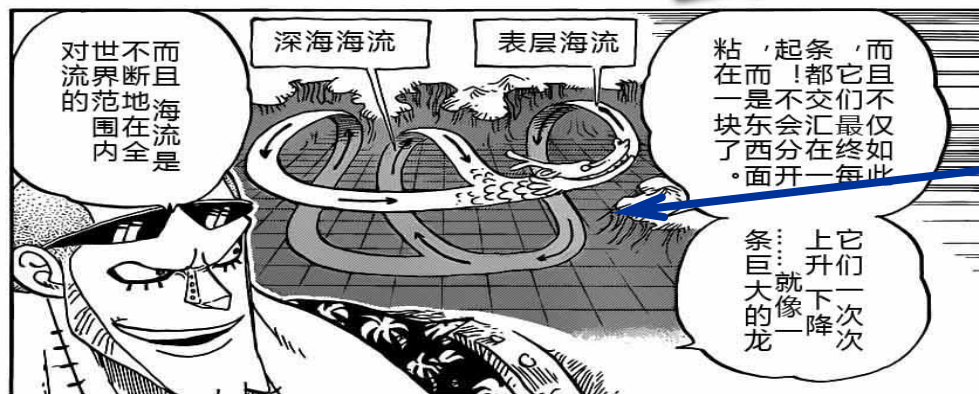
Driven by the sinking of cold, saline water

The heat transport associated with "Atlantic deep circulation" plays an important role in the present climate.

<http://science.howstuffworks.com/ocean-current3.htm>



***Yu-Heng Tseng***



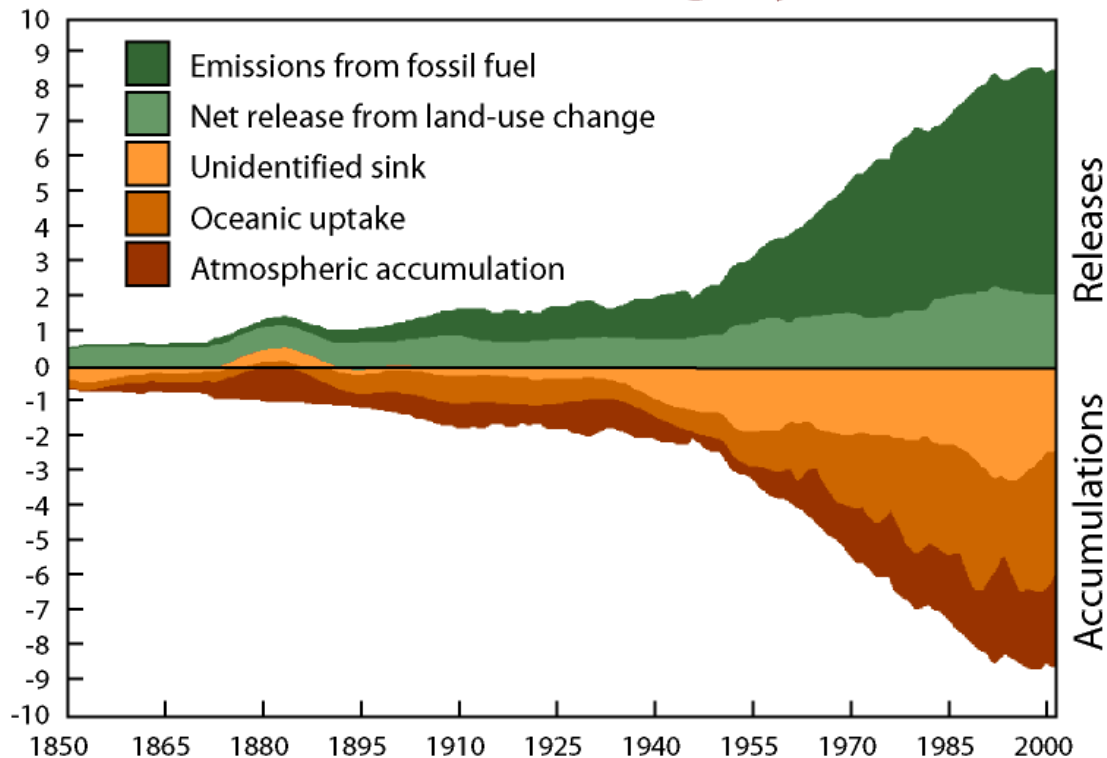
火影海贼死神讨论YY 6370 日英翻、美工招募Q群 12928216

Famous Japanese comic strip:  
海賊王

# Thermohaline circulation

# Global Carbon Budget

## Flux of Carbon (Pg C/yr)



Ocean uptake  
 $\sim 2.4 \text{ PgCyr}^{-1}$   
Mainly biological

A small fraction of  
carbon-flow through  
the ocean system  
So variability can  
have a big effect on  
balance

<http://www.whrc.org/carbon/missingc.htm>

Unidentified sink of  $2.1 \pm 1.1 \text{ PgC/yr}$  needed to balance budget

# Outline

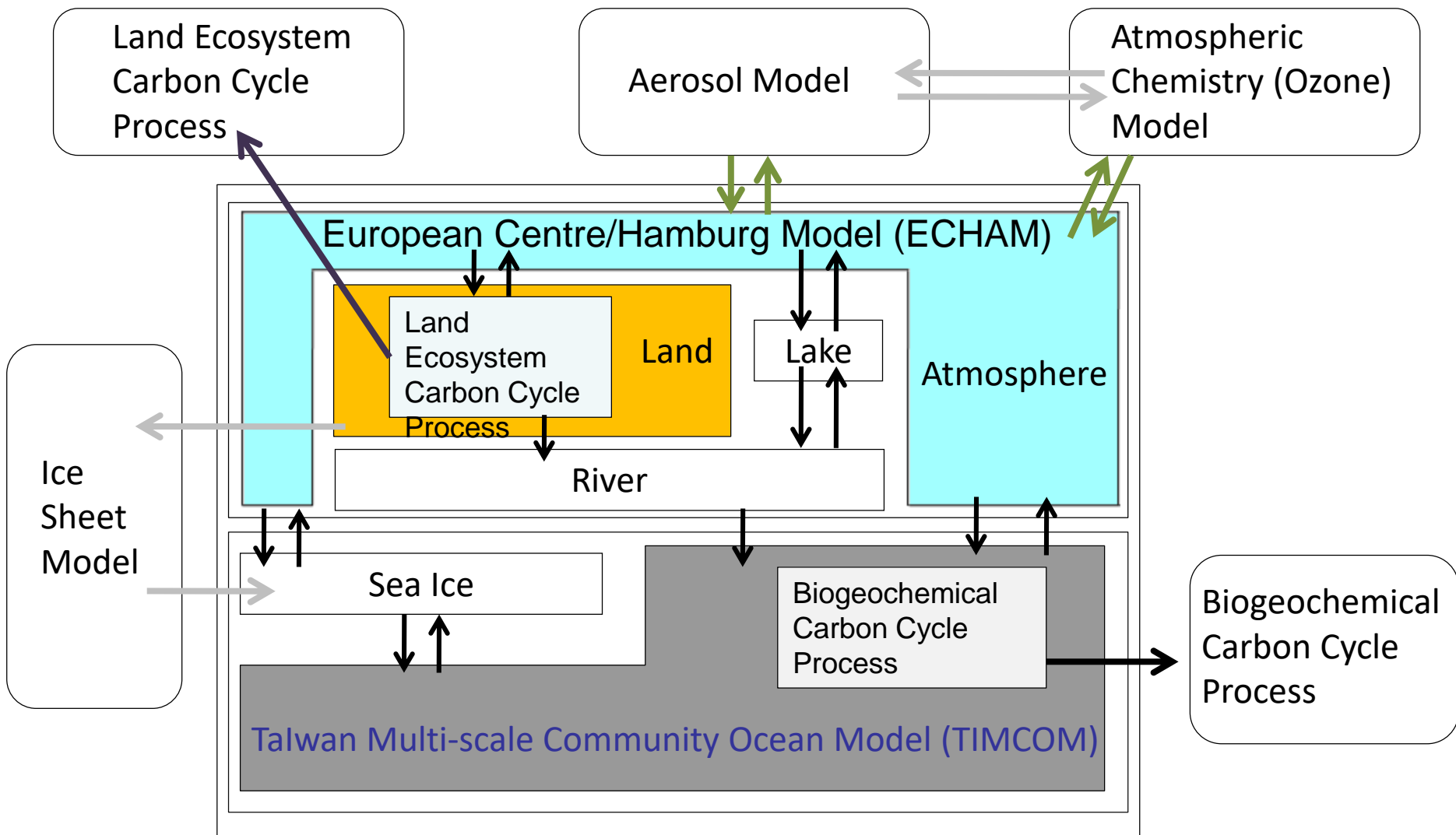
- Objectives and Backgrounds
- Development of TaIwan Multi-scale Community Ocean Model (TIMCOM)
  - <http://140.112.66.144/research/timcom>
  - North Atlantic Ocean Modeling System
  - North Pacific Ocean Modeling System
- Parallel Domain-decomposed TIMCOM (PD-TIMCOM)
  - Parallel EVP solver
  - Preliminary results
- Summary

# Objectives

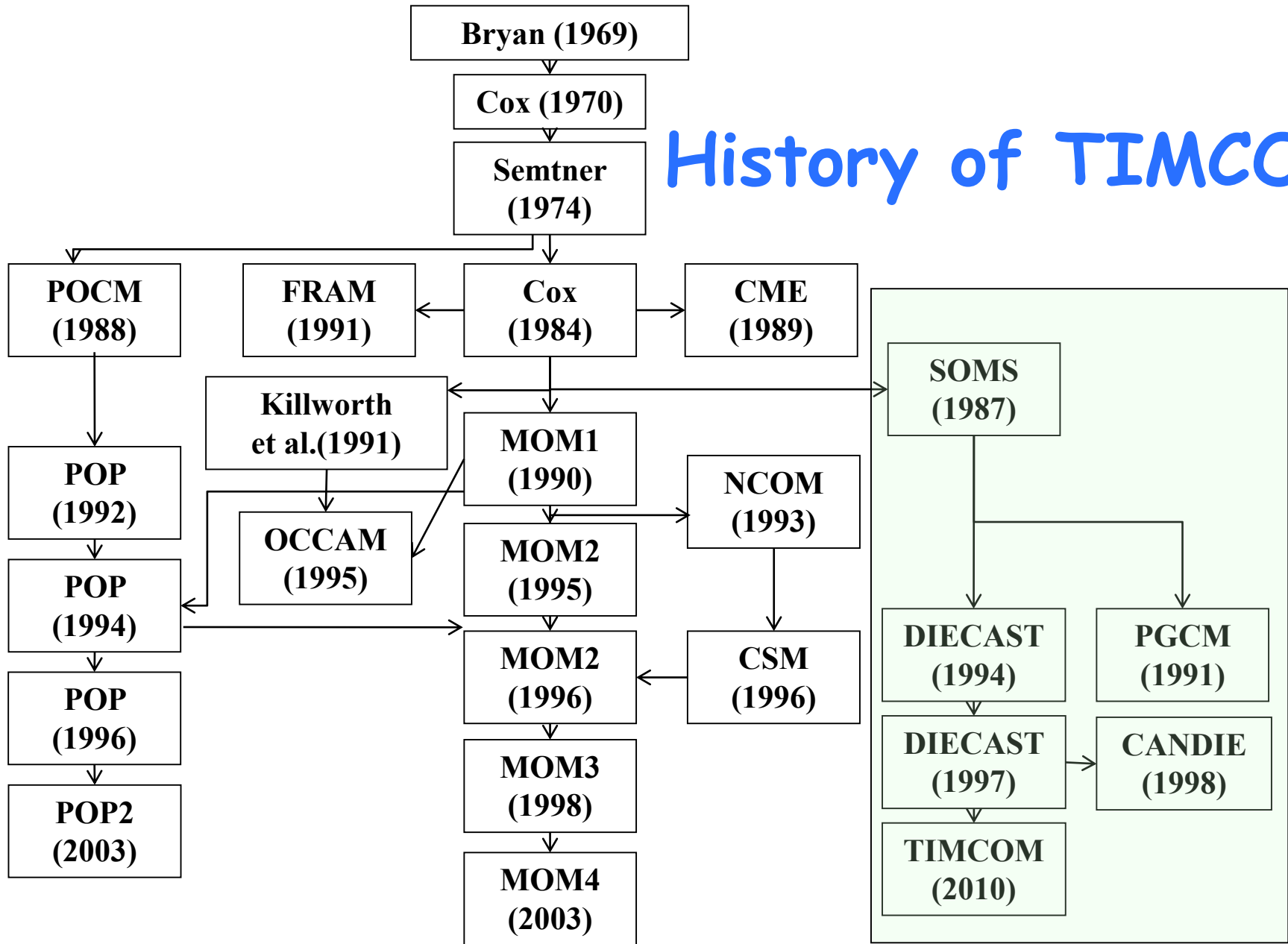
- Building a high performance, multiple-grids global ocean circulation model for global ocean climate study in an **Earth System Model Framework**
- Resolving multi-scale dynamics with the most efficient two-ways coupling approach (serial code) or parallel solver (parallel code) in high accuracy
- Studying the Pacific ocean climate
- Investigating the regional circulation in the vicinity of Taiwan



# Taiwan Earth System Model



# History of TIMCOM



LANL

UK

GFDL

NCAR

TAIWAN

# Governing Equations

$\lambda$ : the longitudinal variable

$\phi$ : the latitudinal variable

$z$ : the vertical variable

Continuity eqn. 
$$\frac{1}{R \cos \phi} \left( \frac{\partial u}{\partial \lambda} + \frac{\partial (v \cos \phi)}{\partial \phi} \right) + \frac{\partial \omega}{\partial z} = 0$$

Momentum eqn.

$$\begin{aligned} \frac{\partial u}{\partial t} &= -\mathcal{L}u + \left( f + \frac{u \tan \phi}{R} \right) v - \frac{1}{\rho_0 R \cos \phi} \frac{\partial p}{\partial \lambda} + D_m u + \frac{\partial}{\partial z} \left( A_u \frac{\partial u}{\partial z} \right) \\ \frac{\partial v}{\partial t} &= -\mathcal{L}v - \left( f + \frac{u \tan \phi}{R} \right) u - \frac{1}{\rho_0 R} \frac{\partial p}{\partial \phi} + D_m v + \frac{\partial}{\partial z} \left( A_v \frac{\partial v}{\partial z} \right) \end{aligned}$$

Conservation eqn. for temperature and salinity

$$\frac{\partial T}{\partial t} = -\mathcal{L}T + D_h T + \frac{\partial}{\partial z} \left( K_T \frac{\partial T}{\partial z} \right)$$

Eqn. of State

$$\rho = \rho(S, T, p)$$

Hydrostatic Eqn. 
$$\frac{\partial p}{\partial z} = -(\rho - \bar{\rho})g$$

$$D_{m(h)} = \frac{A_{m(h)}}{R^2} \left( \frac{1}{\cos^2 \phi} \frac{\partial^2}{\partial \lambda^2} - \tan \phi \frac{\partial}{\partial \phi} + \frac{\partial^2}{\partial \phi^2} \right) \quad \mathcal{L} = \frac{u}{R \cos \phi} \frac{\partial}{\partial \lambda} + \frac{v}{R} \frac{\partial}{\partial \phi} + \omega \frac{\partial}{\partial z}$$

# Outline

- Objectives and Backgrounds
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# MEDiNA model: Bathymetry (km) and sub-domains

Six domain:

GOM ( $1/8^\circ$ )  $304 \times 336$

NAB ( $1/4^\circ$ )  $162 \times 398$

IBE ( $1/8^\circ$ )  $100 \times 794$

VIS ( $1/16^\circ$ )  $60 \times 158$

GIB ( $1/24^\circ$ )  $125 \times 107$

MED ( $1/8^\circ$ )  $316 \times 157$

30 vertical layers;

top layer 11 m thick;

bottom layer 750 m thick

The figure is a map of the Atlantic Ocean region, spanning from 100°W to 40°E and 0° to 60°N. It displays bathymetry with a color scale from -6000 m (dark blue) to 6000 m (dark green). The map is divided into several sub-domains, each outlined with a colored border and labeled with its resolution and dimensions:

- Red border:**  $1/8^\circ \times 1/8^\circ$  (GOM domain)
- Blue border:**  $1/4^\circ \times 1/4^\circ$  (NAB domain)
- Orange border:**  $1/8^\circ \times 1/8^\circ$  (IBE domain)
- Yellow border:**  $1/16^\circ \times 1/16^\circ$  (VIS domain)
- White border:**  $1/24^\circ \times 1/24^\circ$  (GIB domain)
- Black border:**  $1/8^\circ \times 1/8^\circ$  (MED domain)

Geographical labels include "Greenland" and "Ireland". The map includes latitude and longitude markings along the axes and a color bar at the bottom indicating bathymetry in meters (m).

Yu-Heng Tseng

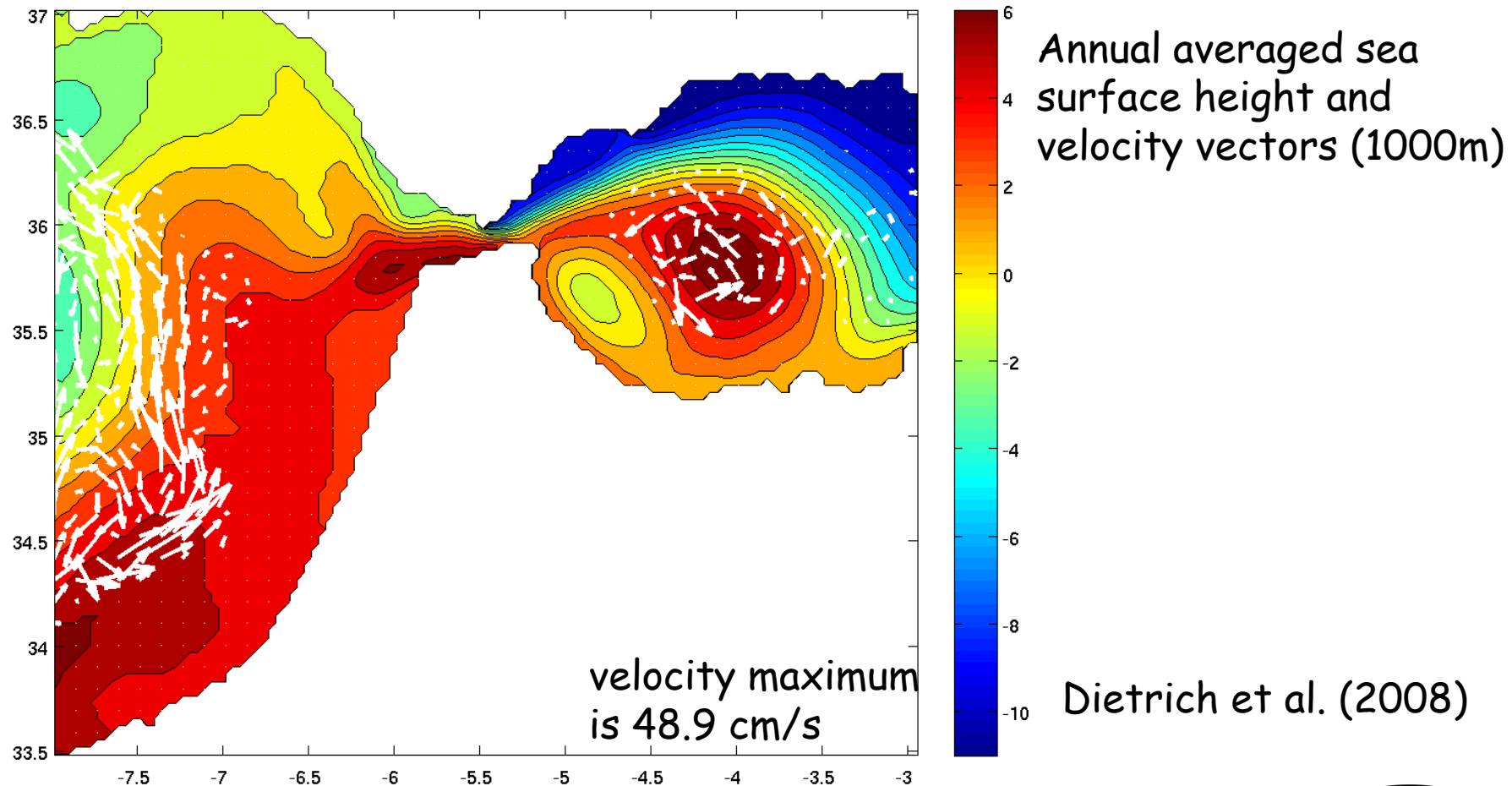
HC/EFDL, NTU

HIGH-PERFORMANCE COMPUTING & ENVIRONMENTAL FLUID DYNAMICS LABORATORY  
National Taiwan University  
Atmospheric Sciences

- 4<sup>th</sup> order accurate, Z-level, rigid-lid approximation
- Mixed Arakawa "a" and "c" grid TaiCOM
- The control volume equations include fluxes of the conservation of momentum, heat and salt across control volume faces.
- Bathymetry:
  - Interpolated from unfiltered ETOPO2 depth data
  - Supplemented with NCOR's 1-minute high accuracy depth archive.
- The vertical resolution ~ linear-exponential stretched grid, 26 layers
  - Z=6, 20, 36, 54, 75, 98, 126, 159, 198, 244, 298, 364, 442, 537, 652, 790, 958, 1161, 1408, 1709, 2075, 2520, 3063, 3725, 4532 m.
- Varying latitude and uniform longitude grid (Mercator grid).
- Surface forcing:
  - Use interpolated monthly Hellerman and Rosentstein winds (Hellerman and Rosenstein, 1983).
  - Use Levitus'94 climatology (Levitus and Boyer, 1994) to initialize the model and determine its surface sources of heat and fresh water.
- The northern boundary is closed. The southern boundary condition (30° S) is slow nudging toward climatology in a sponge layer. The bottom is insulated, with non-slip conditions parameterized by a nonlinear bottom drag.
- Sub-grid scale vertical mixing is parameterized by eddy diffusivity (for temperature and salinity) and viscosity (for momentum) using a modified Richardson number based approach based on Pacanowski and Philander (1982)

# Theory, observation and modelling

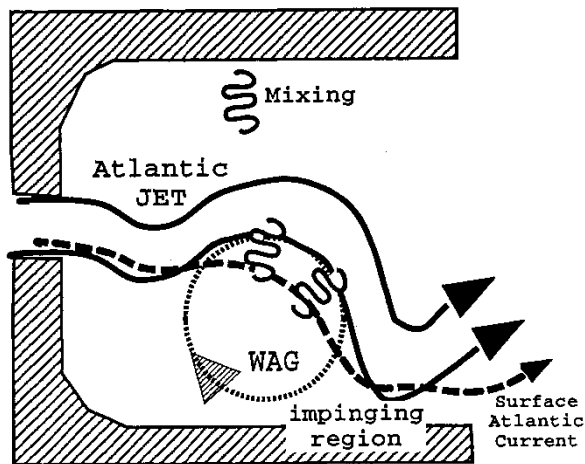
## Strait of Gibraltar domain



# Theory, observation and modelling

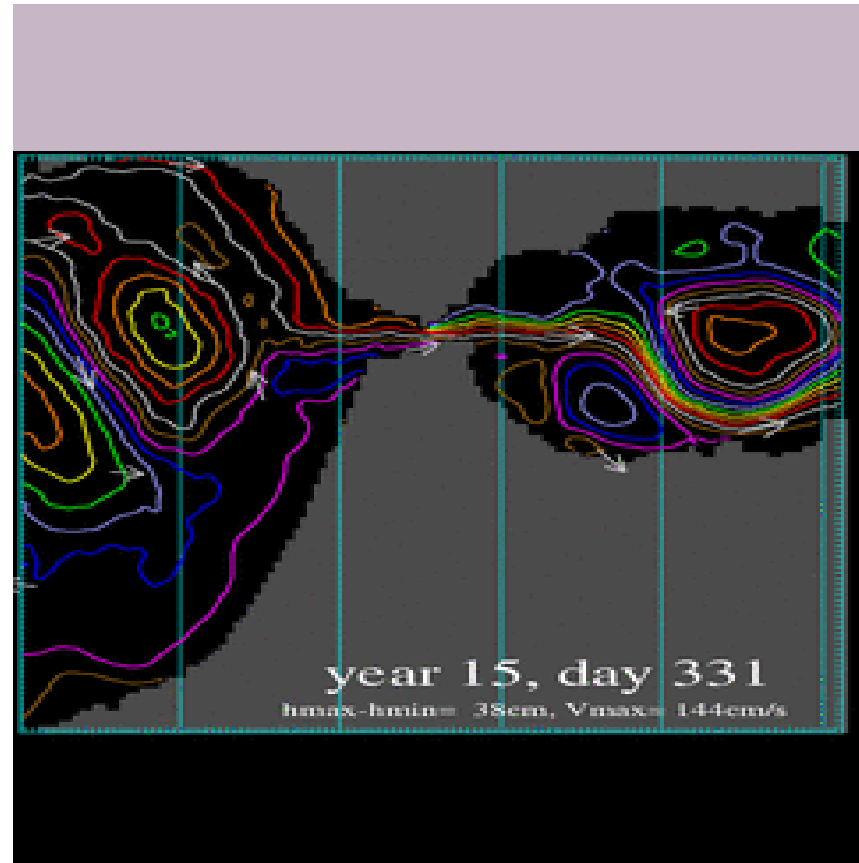
## Strait of Gibraltar domain

### Western Alboran Gyre (WAG)



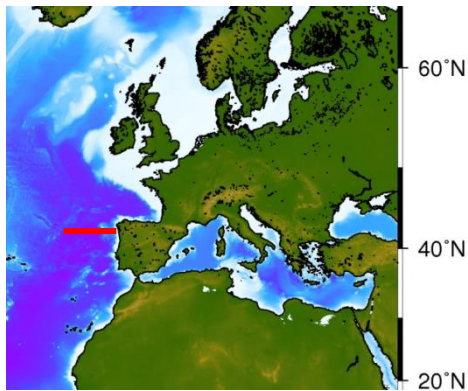
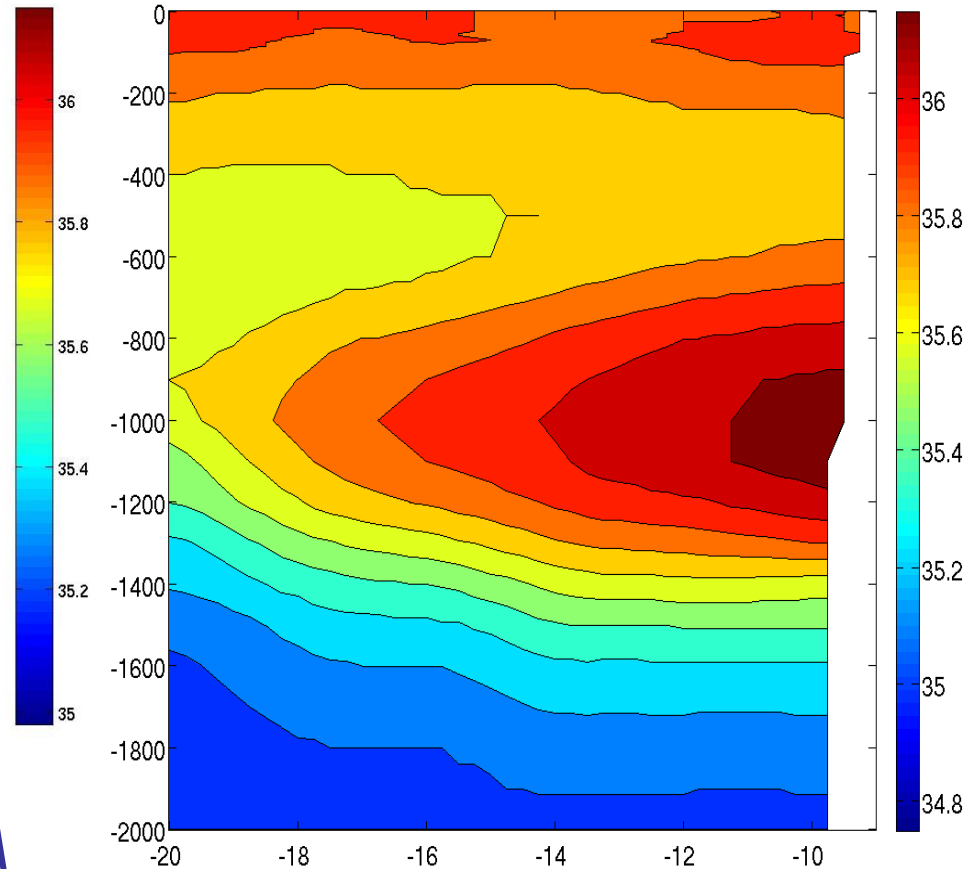
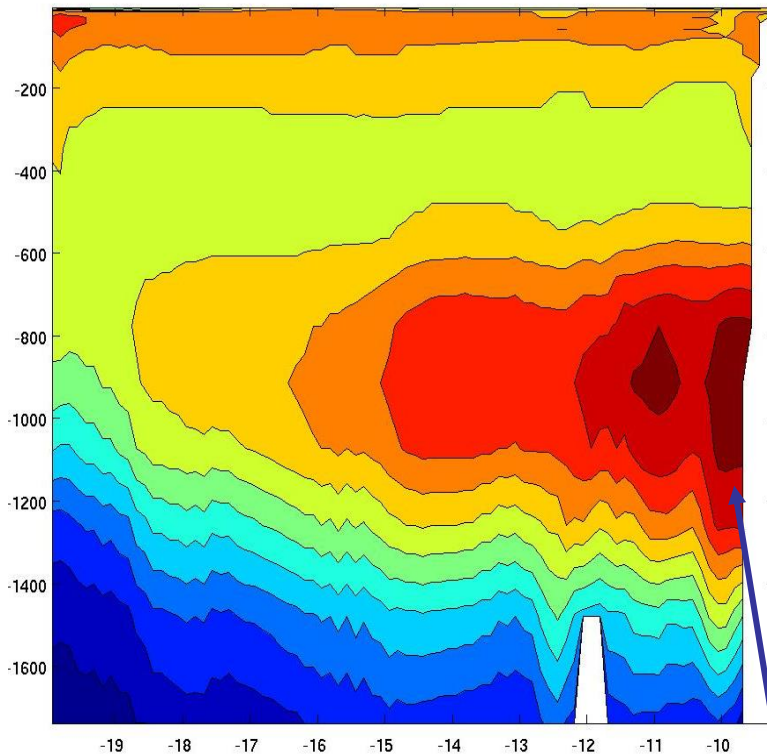
**Figure 3.** Sketch of the upper circulation (0-200 m) in the western Alboran basin. The surface Atlantic current, characterized by a salinity minimum, may enter into the western Alboran gyre (WAG) crossing the iso-lines of dynamic height anomaly produced by the deeper density gradients of the gyre. After mixing briefly with water in the core of the WAG, this surface Atlantic current leaves the WAG through the impinging region on the African coast.

**Viúdez, Pinot and Haney (1998)**





# Vertical/longitudinal salinity section at 43°N



Annual averaged  
model results

U. S. Navy's GDEM climatology

High salinity water affecting the global THC

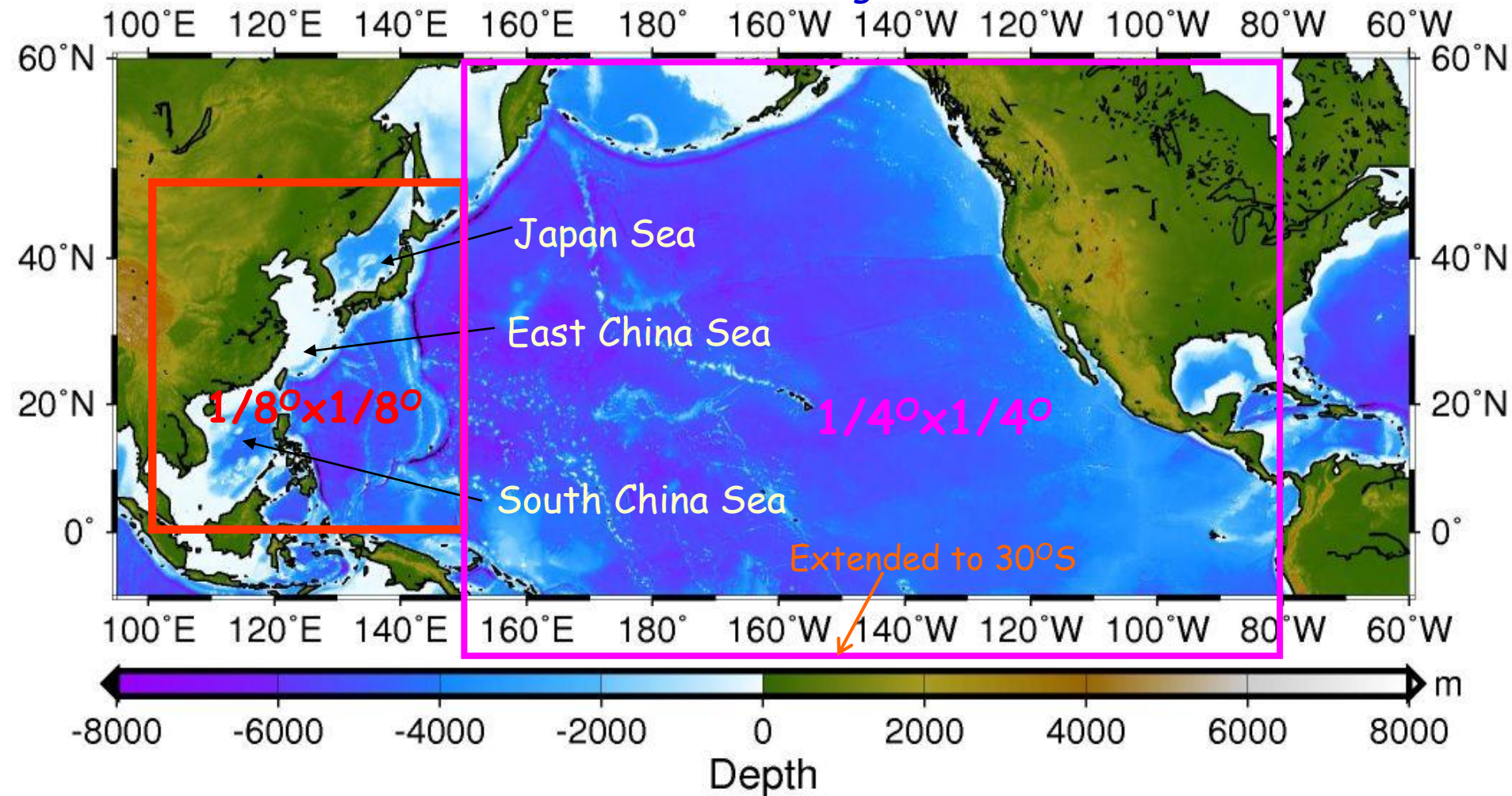
# Outline

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# Dual-grid Pacific Ocean Model (DUPOM)

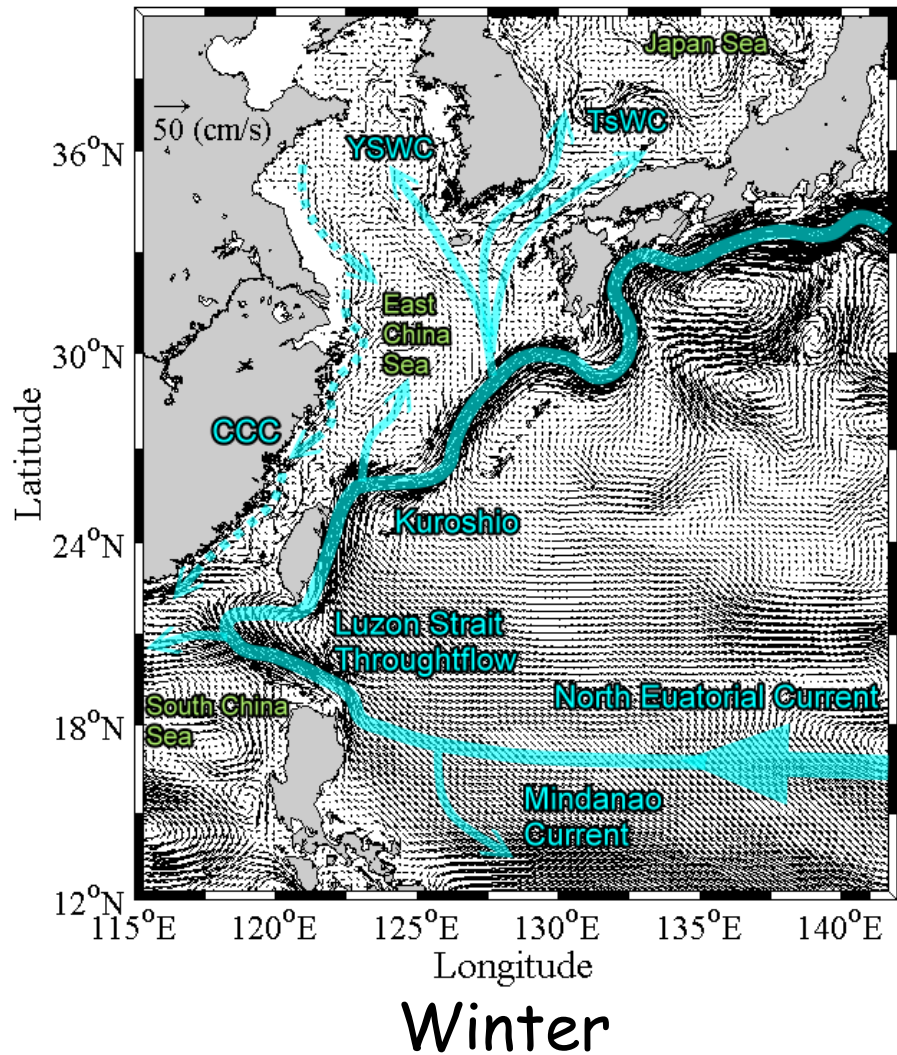
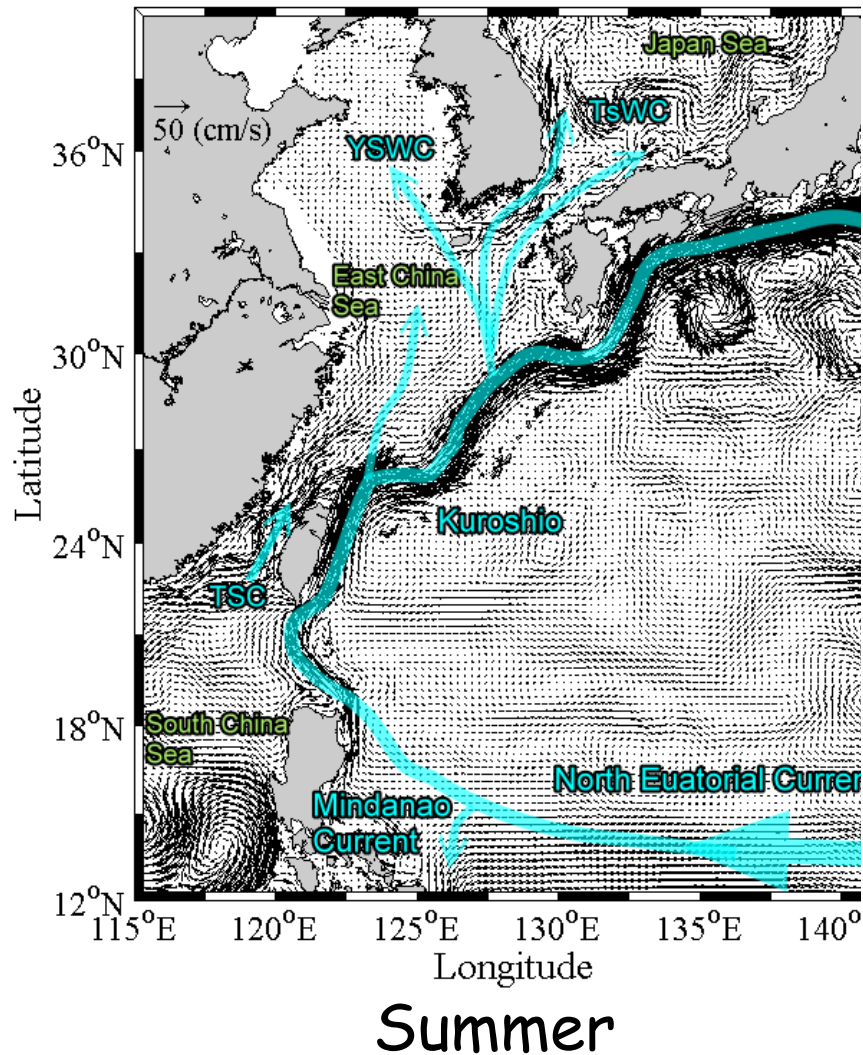
**NPB (North Pacific Basin) Domain:**  
30°S to 60°N and 150°E to 80°W  
grid resolution 1/4°

**TAI Domain:**  
0° to 50°N and 100°E to 150°E  
grid resolution 1/8°





# General seasonal circulation pattern-DUPOM





# Outline

- Objectives and Backgrounds
- Development of TaIwan Multi-scale Community Ocean Model (TIMCOM)
  - North Atlantic Ocean Modeling System
  - North Pacific Ocean Modeling System
- New Parallel Domain-decomposed TIMCOM (PD-TIMCOM)
  - Parallel Error Vector Propagation (EVP) solver (Tseng and Chien, 2011, C&F)
  - Preliminary results
- Summary

- To solve the pressure correction equation in OGCMs with hydrostatic approach, an efficient Poisson equation solver is needed.

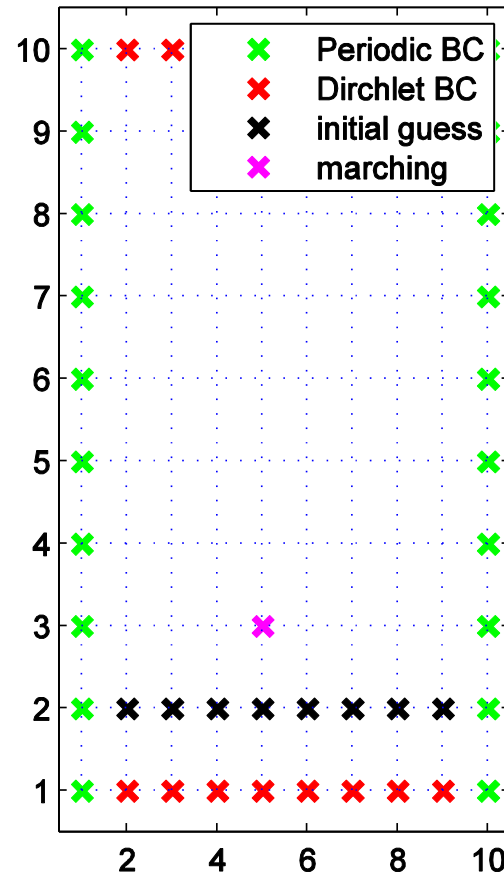
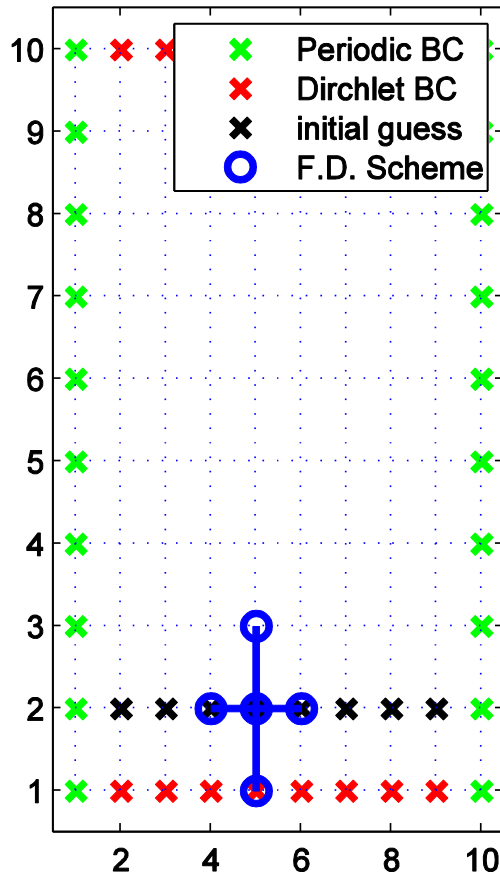
$$\begin{cases} \frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2} = \Phi \\ \text{Boundary Condition} \end{cases}$$

- In a finite difference discretization form

$$\frac{u_{i+1,j} - 2u_{i,j} + u_{i-1,j}}{\delta x^2} + \frac{u_{i,j+1} - 2u_{i,j} + u_{i,j-1}}{\delta y^2} = \phi_{i,j}$$

- Choose an initial guess  $u'_{i,2}$  from the boundary, where  $u_{i,2} = u'_{i,2} + e$
- March the solution over the whole domain.

$$u'_{i,j+1} = \delta y^2 \phi_{i,j} + 2u'_{i,j} \left( 1 + \frac{\delta y^2}{\delta x^2} \right) - \frac{\delta y^2}{\delta x^2} (u'_{i-1,j} + u'_{i+1,j}) - u'_{i,j-1}$$

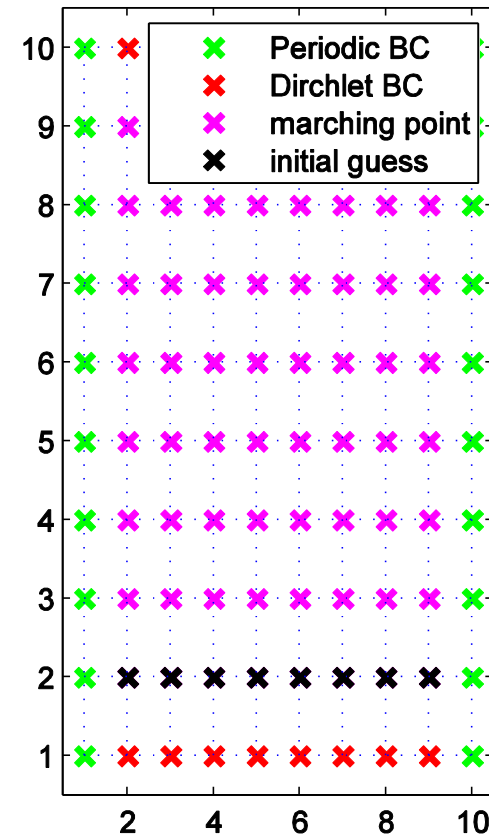
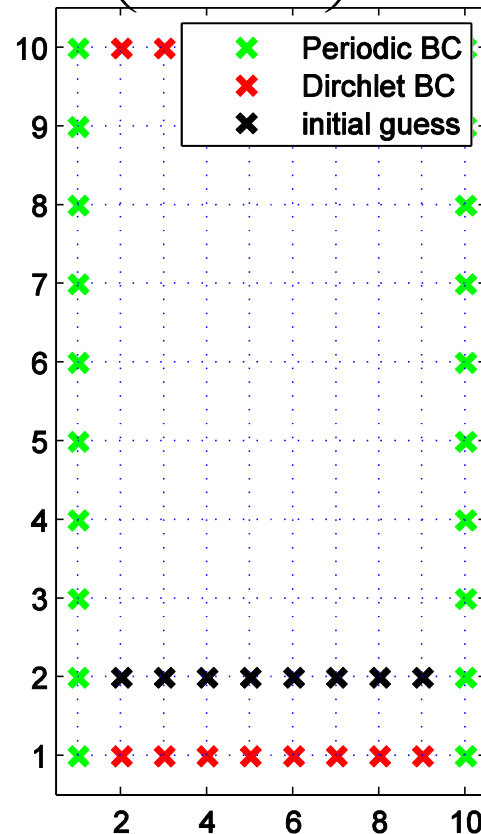


- To establish the linear relation between vector along  $e_{i,2}$  in terms of  $e_{i,J0}$ . We obtain the recursion relation for the error propagation.

$$e_{i,j+1} = 2e_{i,j} \left( 1 + \frac{\delta y^2}{\delta x^2} \right) - \frac{\delta y^2}{\delta x^2} (e_{i-1,j} + e_{i+1,j}) - e_{i,j-1}$$

- With this formula, we can establish the linear system of  $e_{i,2}$  and  $e_{i,J0}$ .

$$F_l = C_{lm} E_m$$

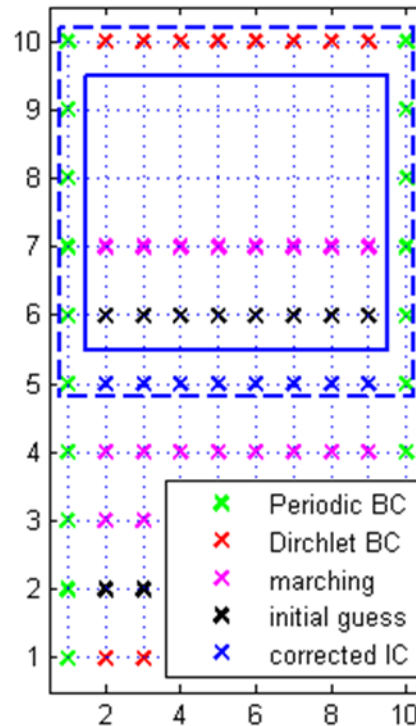
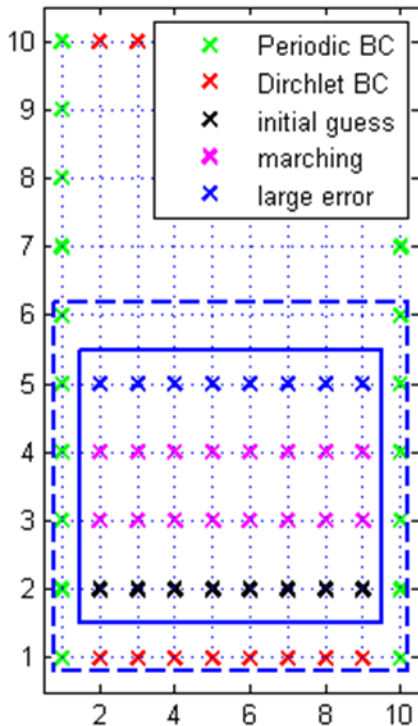




# Stabilized Error Vector Propagation Method

- However, EVP will failed because influence matrix can't afford the round-off error.

$$R = \frac{u'_{i+1,j} - 2u'_{i,j} + u'_{i-1,j}}{\delta x_i^2} + \frac{u'_{i,j+1} - 2u'_{i,j} + u'_{i,j-1}}{\delta y^2} - \phi_{i,j}$$

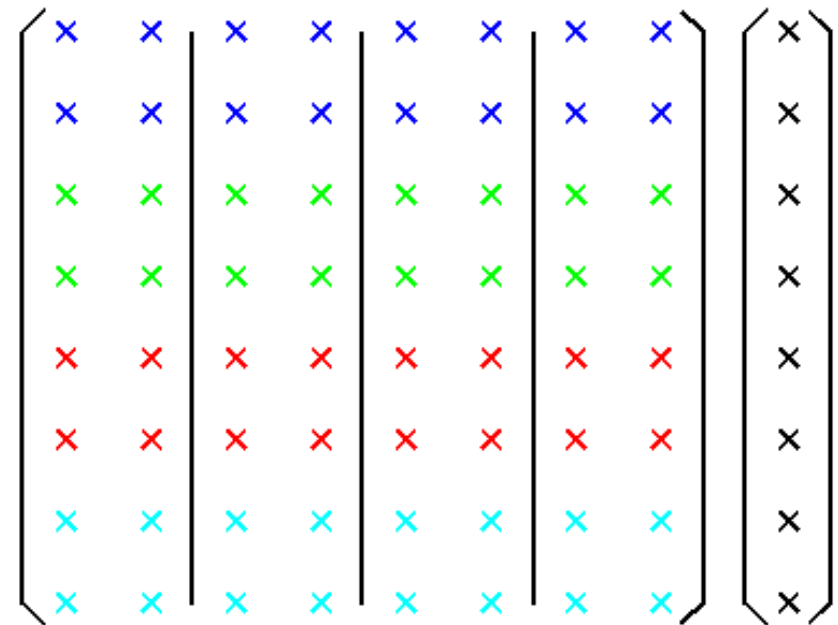
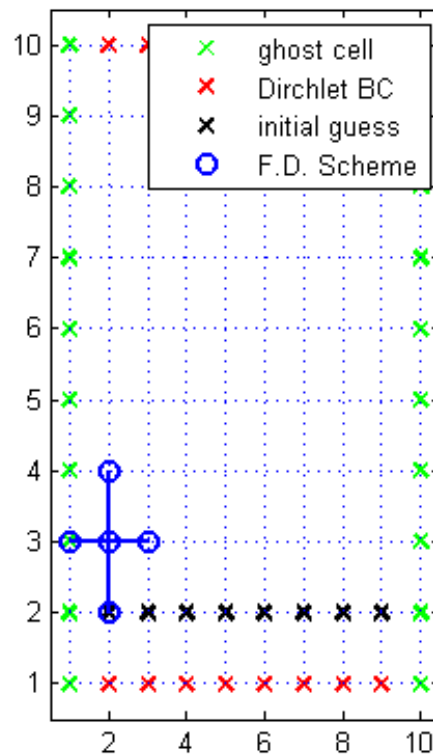
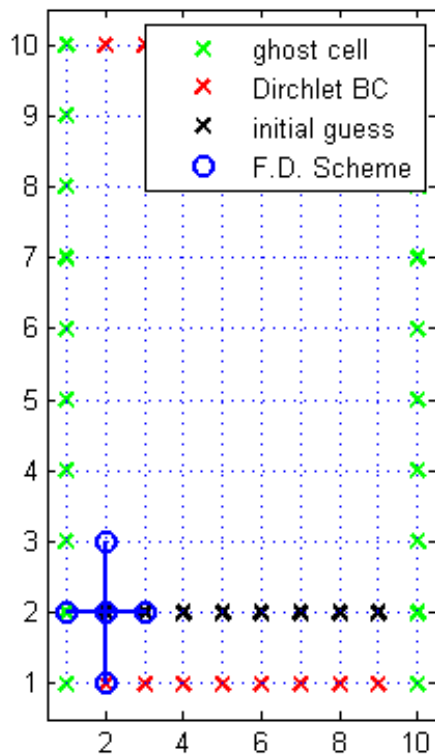


We may apply the residual at blue point, take the influence matrix product and march again to get corrected interior initial condition (IC).

The error of EVP will not blow up enough block along  $y$  direction is applied.

# Parallel Implement to SEVP

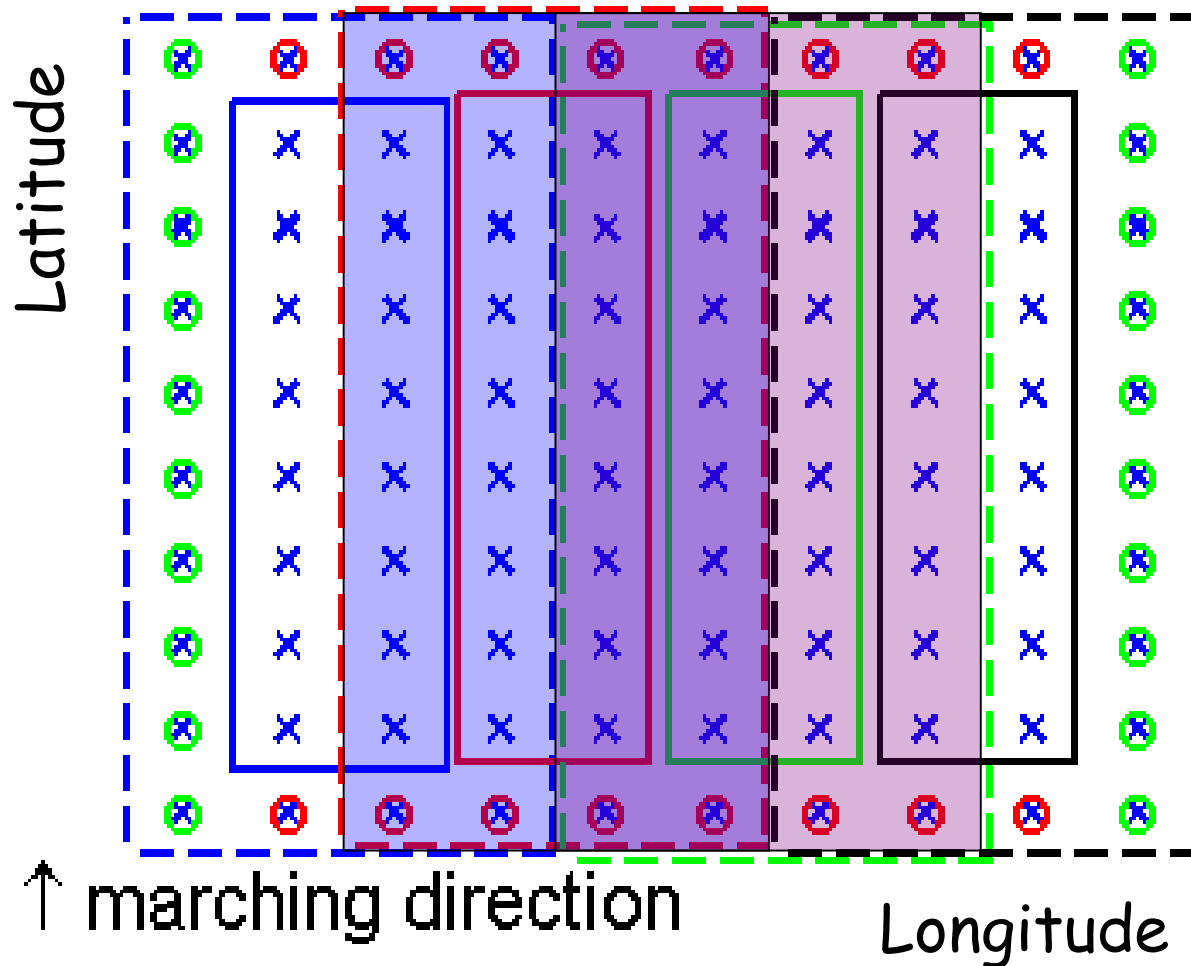
- Requirement of differential operator.
- Use of ghost-cell to compete the scheme.



Decomposed matrix product, colorful point denotes the vertical processor and solid line denote the horizontal processor.

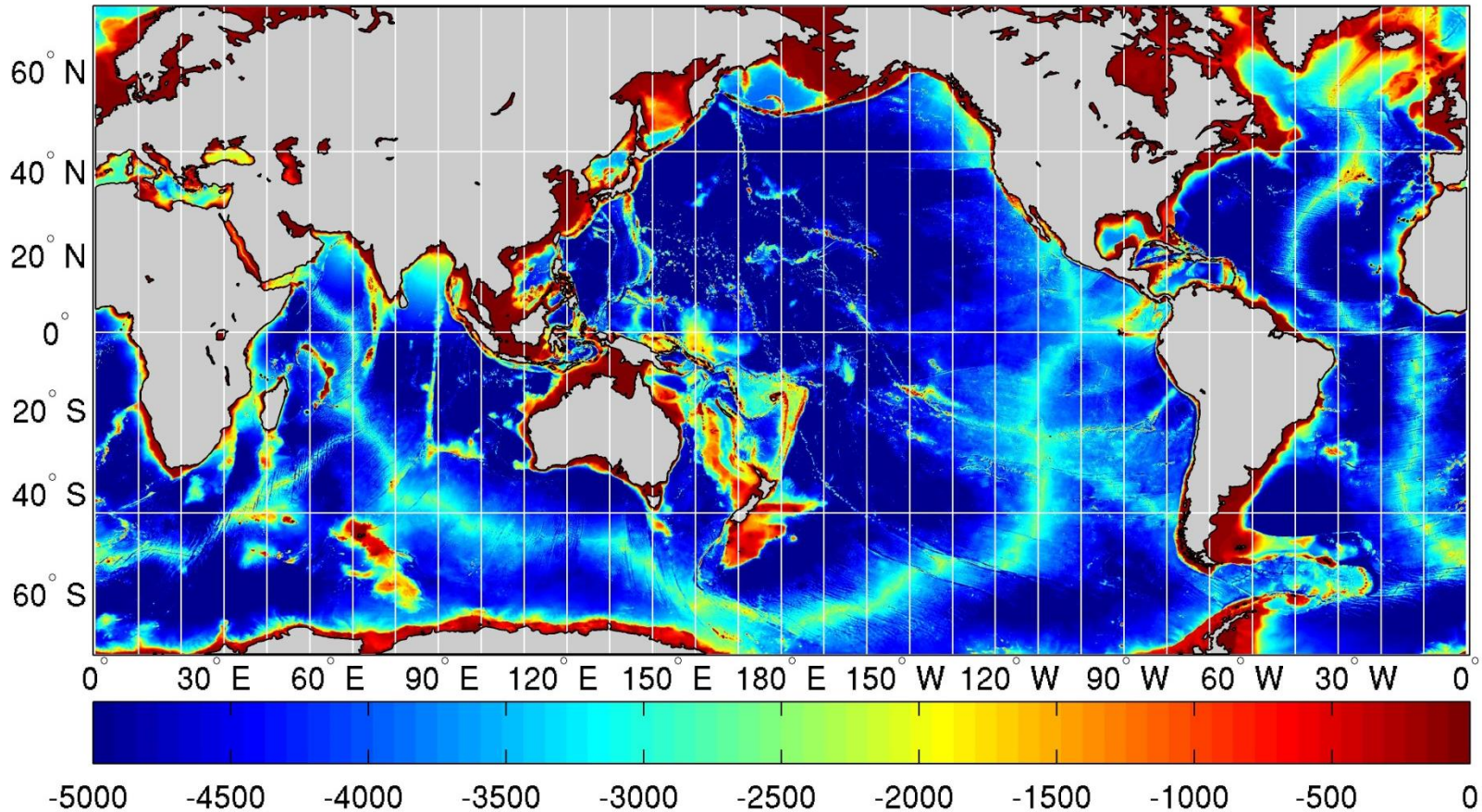
# Parallel overlapping and marching

-Parallel Error Vector Propagation Method (Parallel EVP)



# Example: 192(16x12) decomposed domains

Model Bathymetry and Domain Decomposition



# Platforms for parallel performance test

Platform	CPU	Total cores	Network
IBM C1350	Woodcrest 3.0GHz (dual-core x2)	4096	InfiniBand
Cray XT5	Opteron 2.4GHz (quad-core x2)	5312	SeaStar
IBM iData Plex	Intel Nehalem 2.67GHz (quad-core x2)	3200	InfiniBand



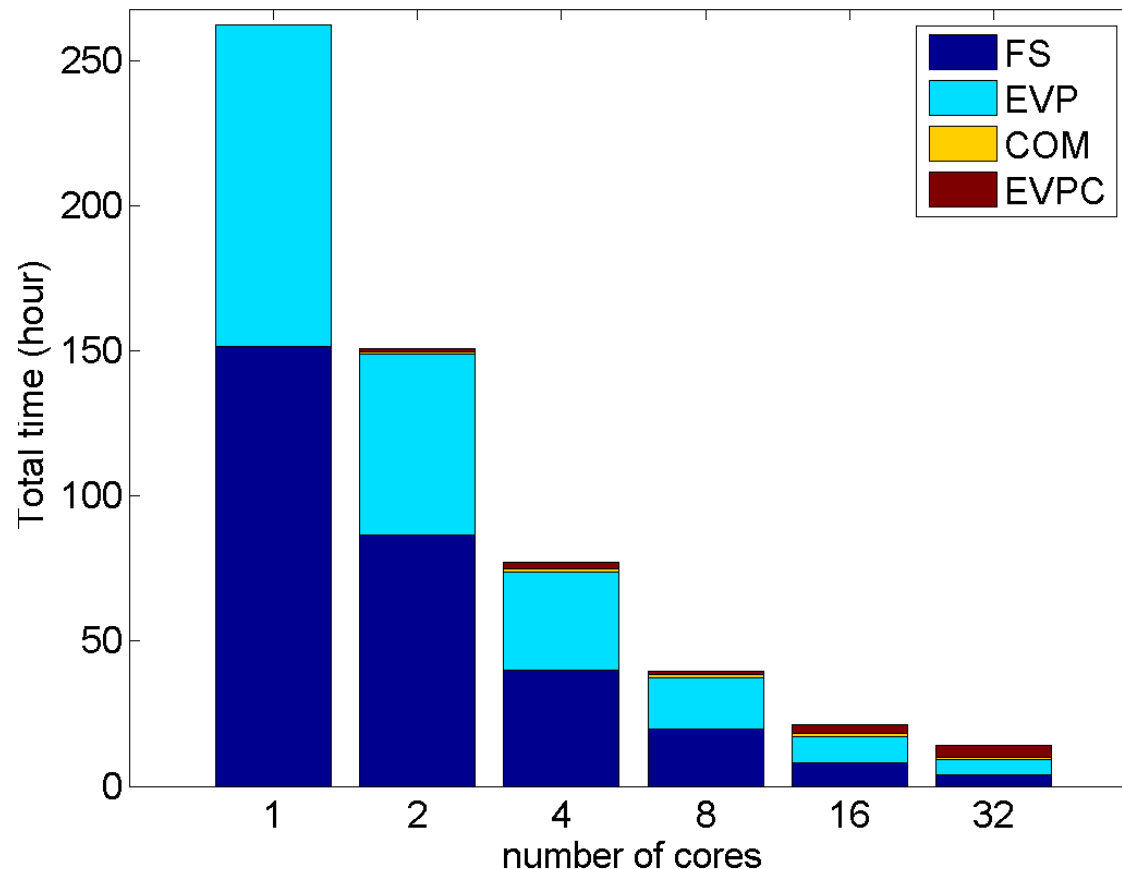
# Parallel performance

FS: simulation time

EVP: pressure solver time

COM: Other Communication

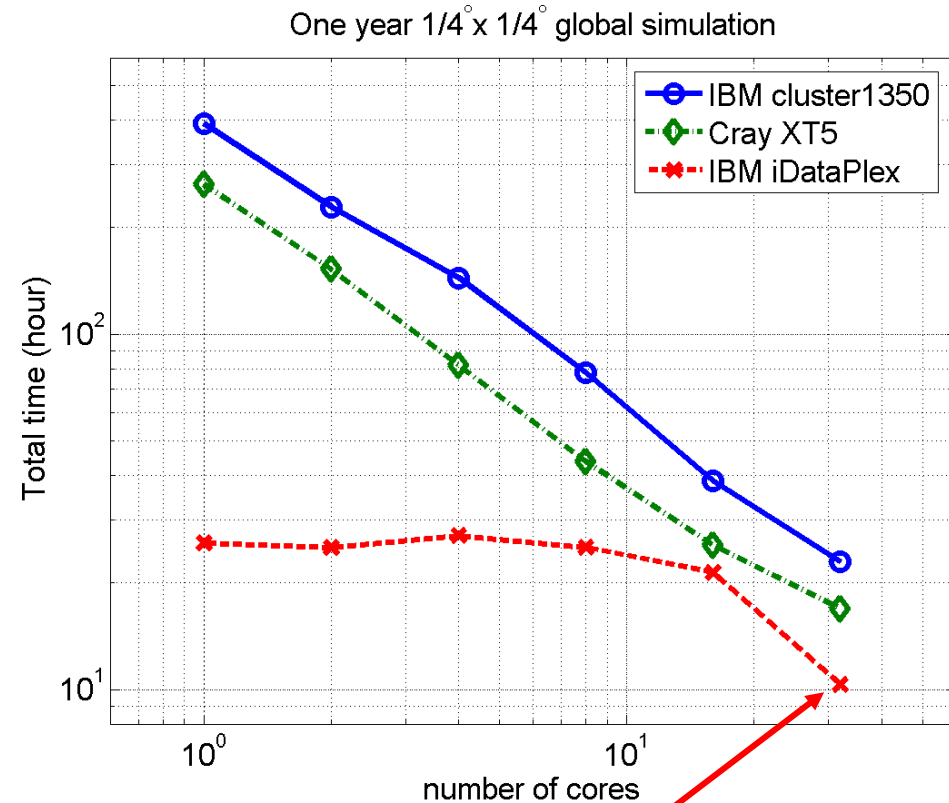
EVPC: Communication in EVP solver



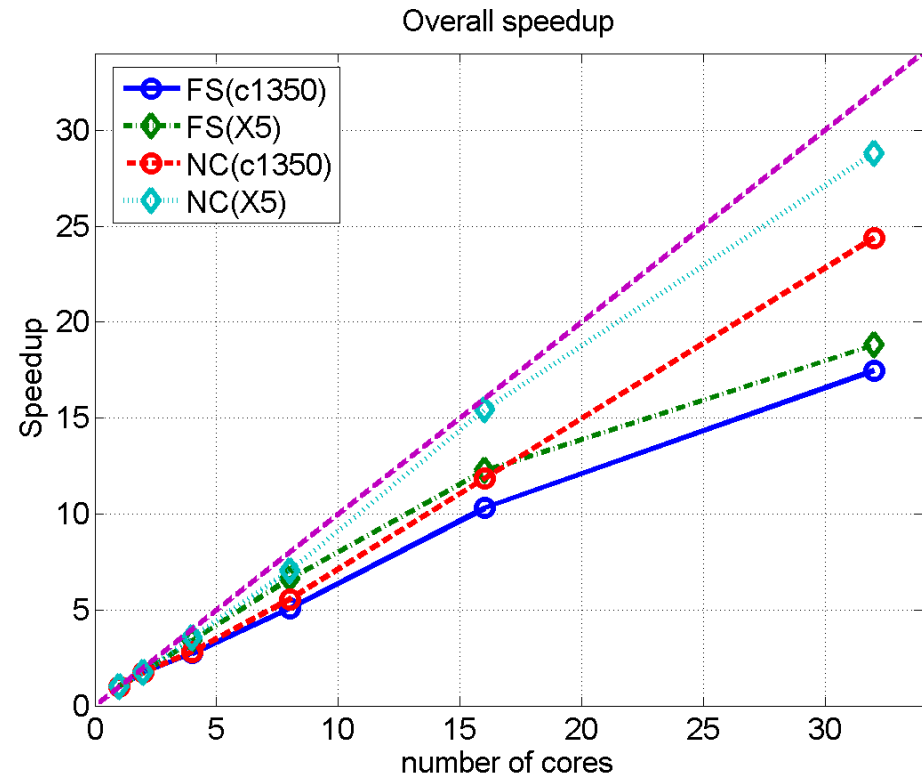
# Parallel speedup

FS: all computation

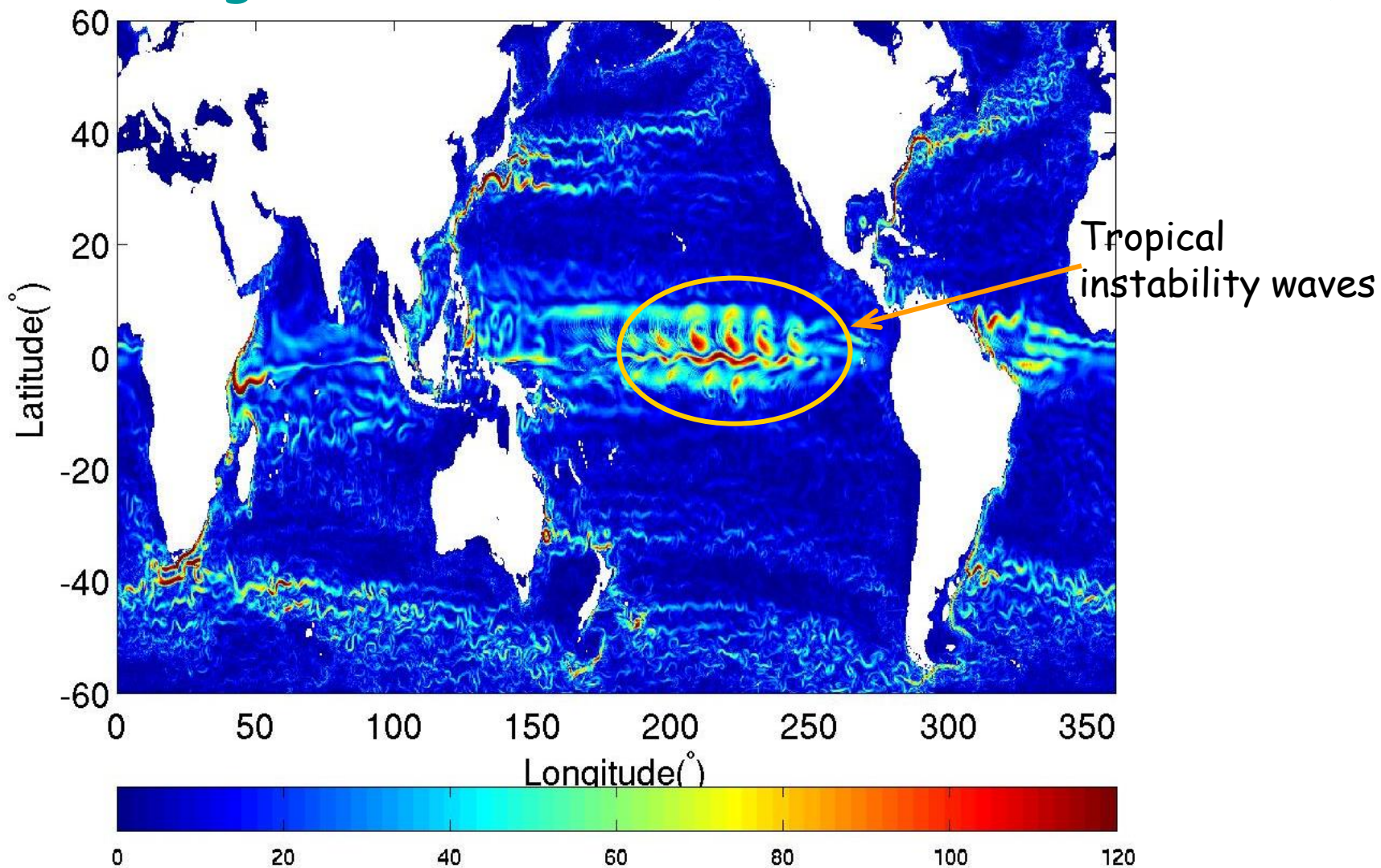
NC: full computation without the communication



Roughly 10 CPU hrs/per simulation year



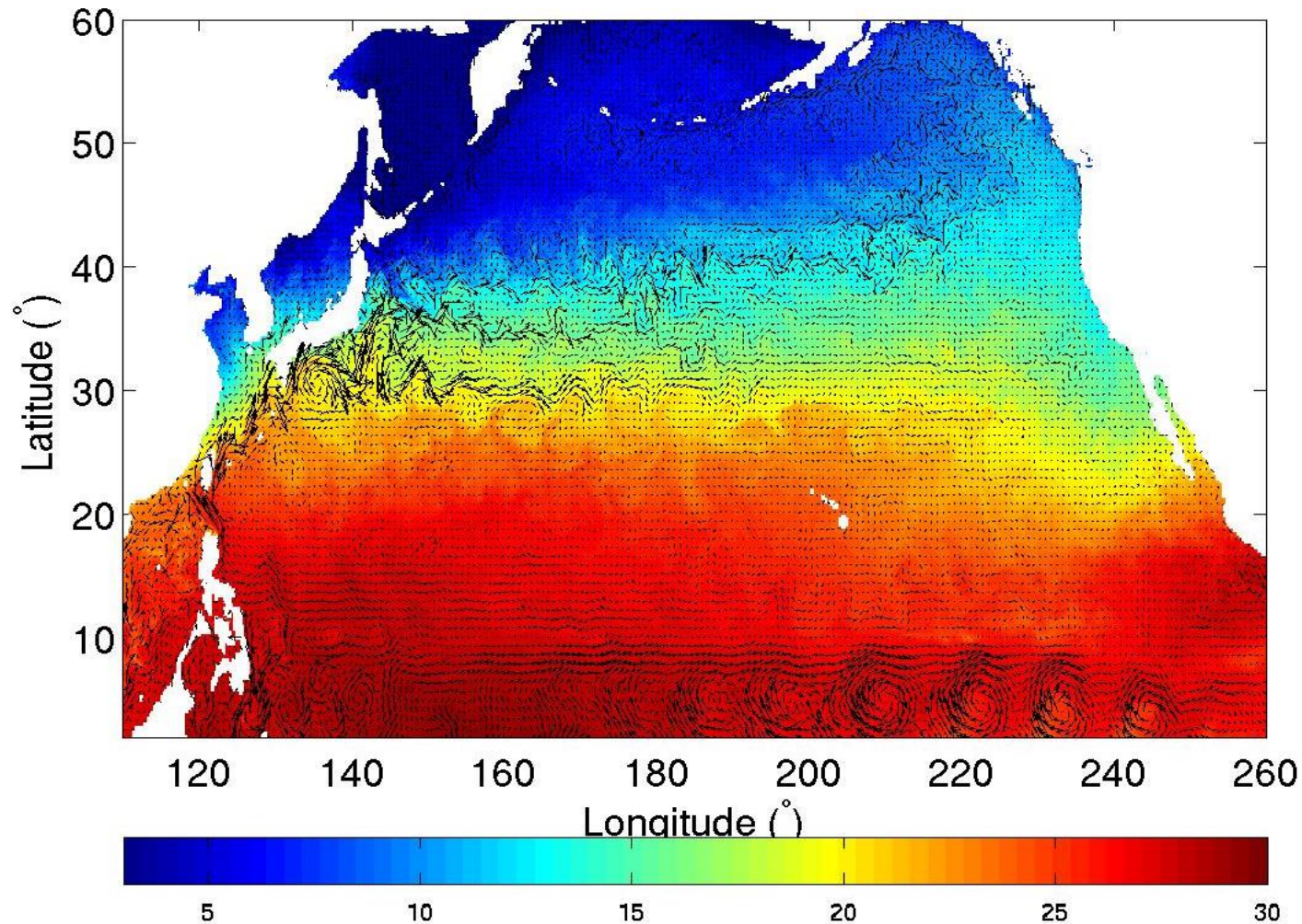
$1/4^\circ \times 1/4^\circ$  global resolution (domain  $1442 \times 720 \times 26$ )



Global velocity speed field (day 5, Year 49)

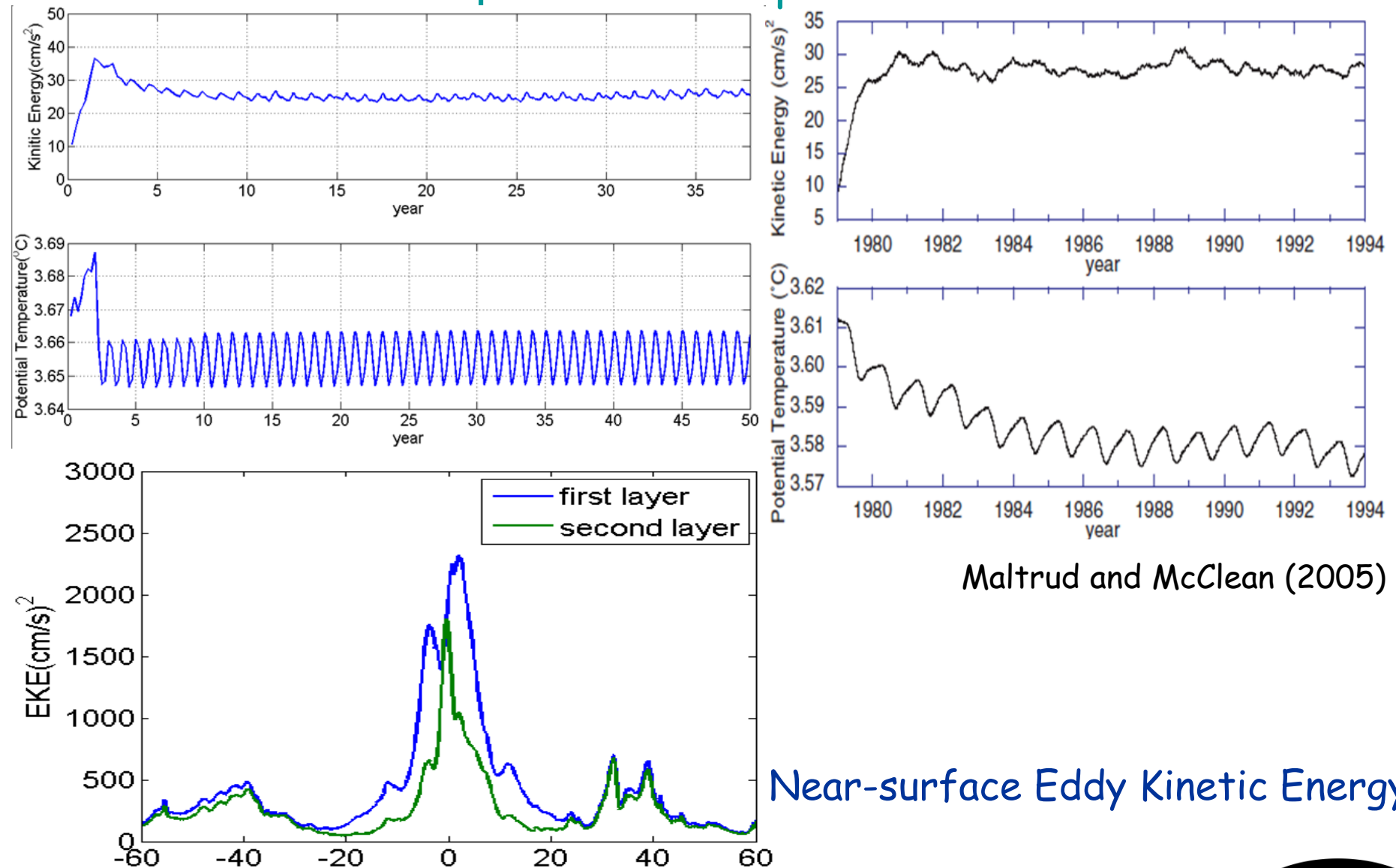


$1/4^\circ \times 1/4^\circ$  global resolution (domain  $1442 \times 720 \times 26$ )



North Pacific temperature and velocity field (day 5, Year 49)

# Time evolution of globally averaged Total Kinetic Energy and potential temperature



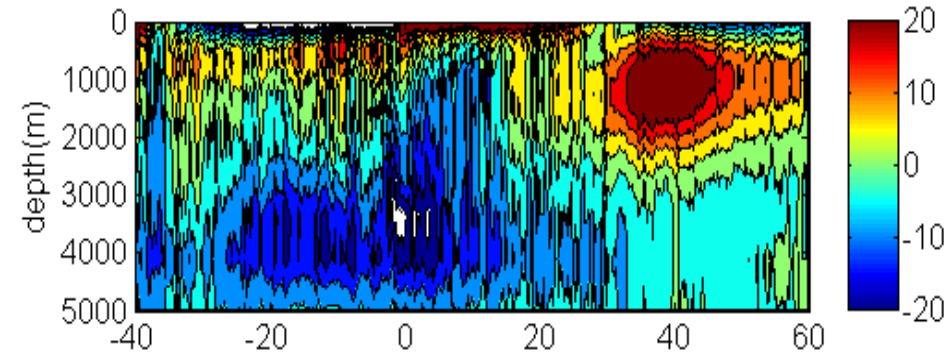
Maltrud and McClean (2005)

Near-surface Eddy Kinetic Energy

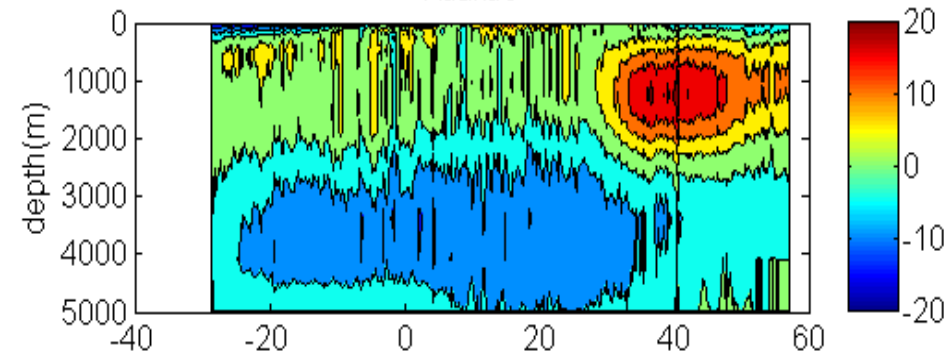


# Meridional overturning streamfunction and global heat transport

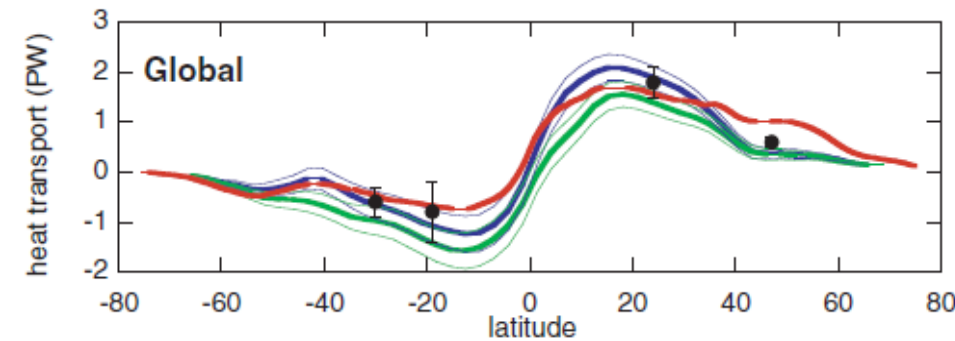
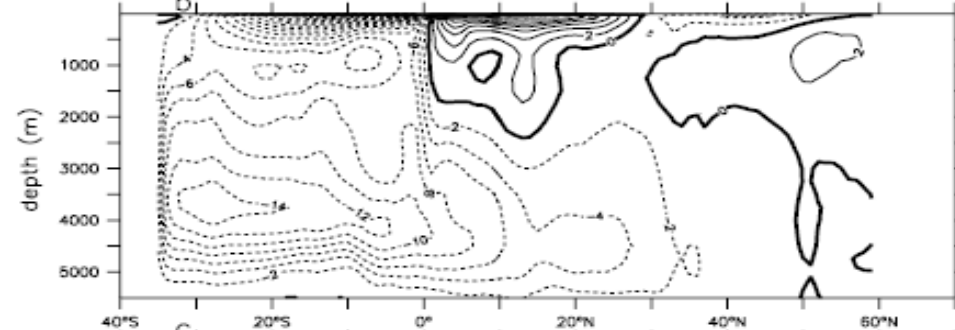
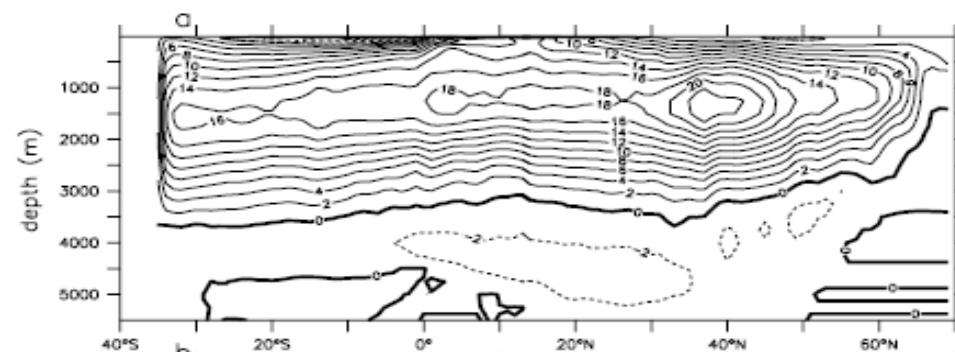
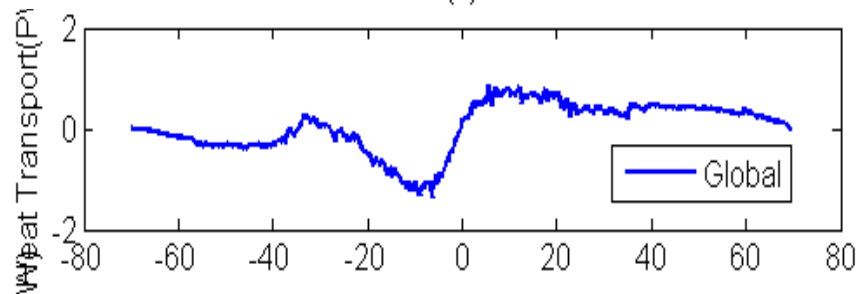
Global



Atlantic

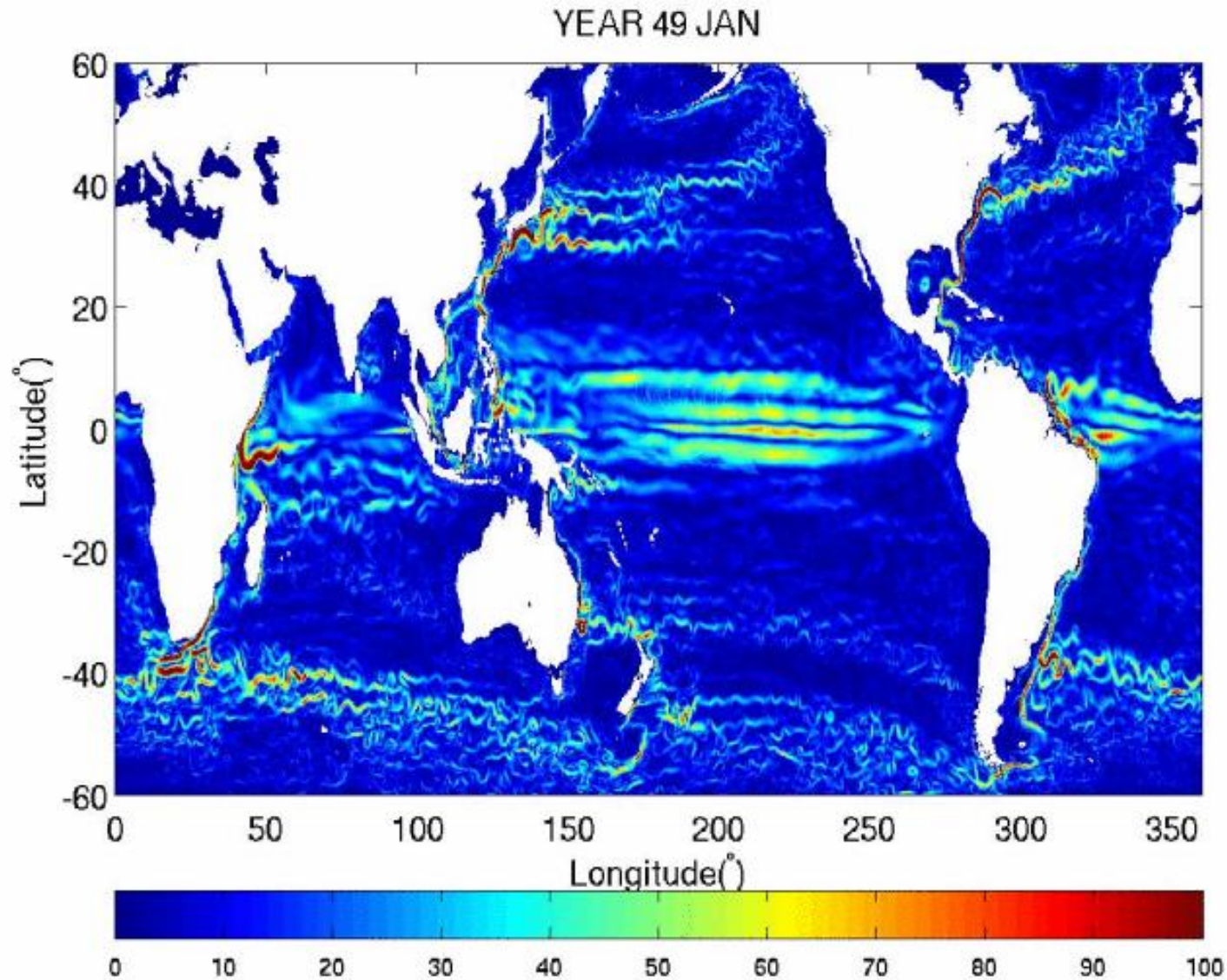


Latitude( $^{\circ}$ )



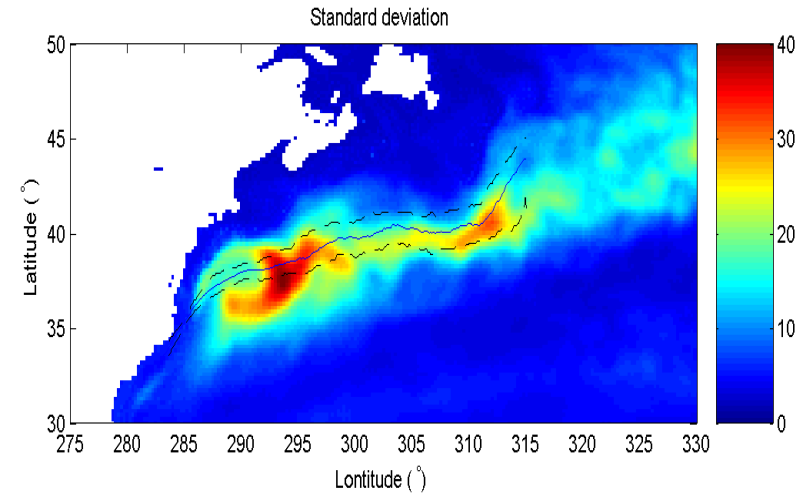
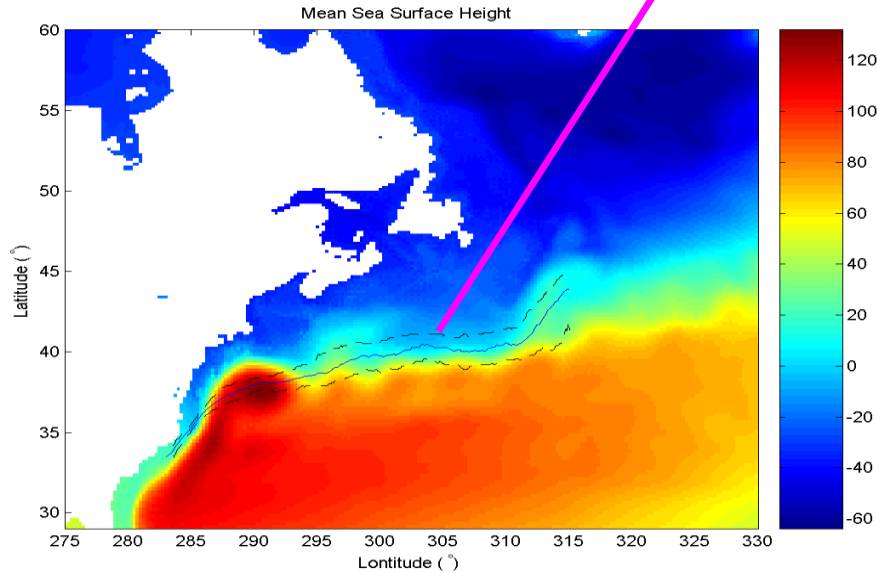
Maltrud and McClean (2005)

# Animation of global surface velocity speed

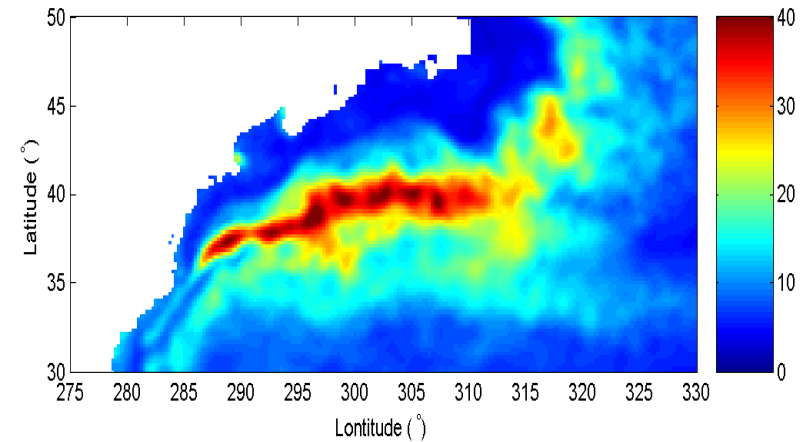
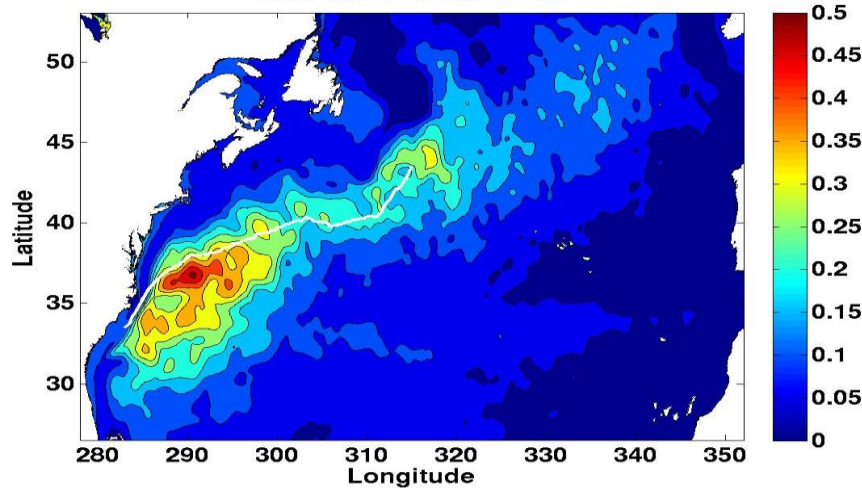


# Gulf Stream

The mean Gulf Stream IR northwall pathway  $\pm 1\sigma$  (standard deviation) by Cornillon and Sirkes



HYCOM Run 9.4 (12-15) Sea Surface Height Standard Deviation with NAVO IR Mean Path

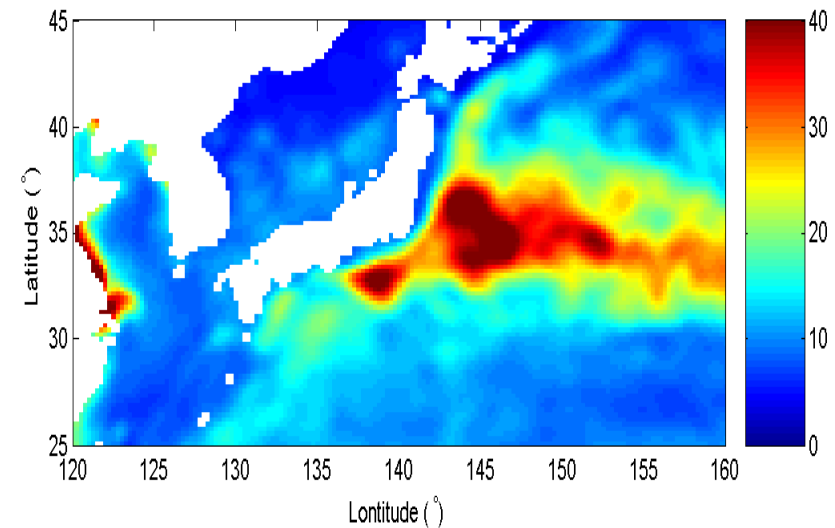
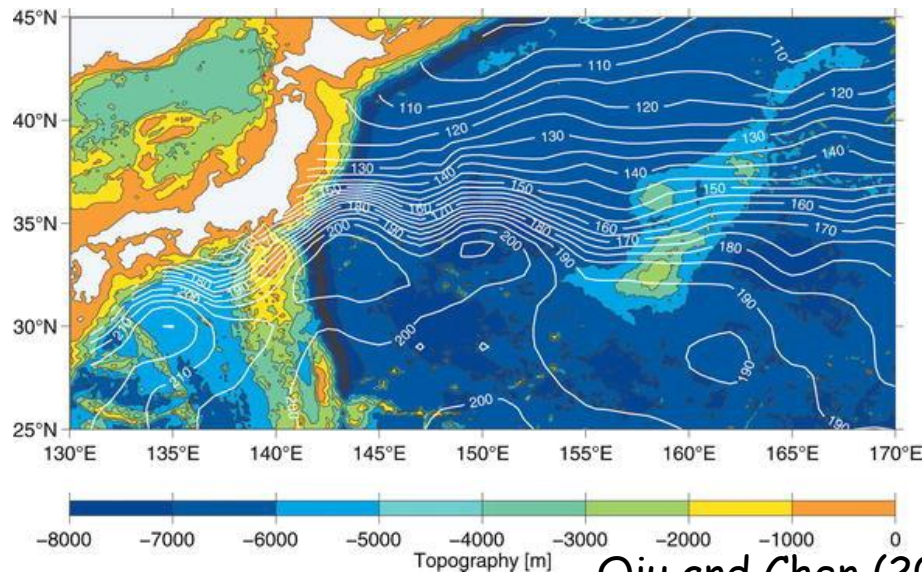
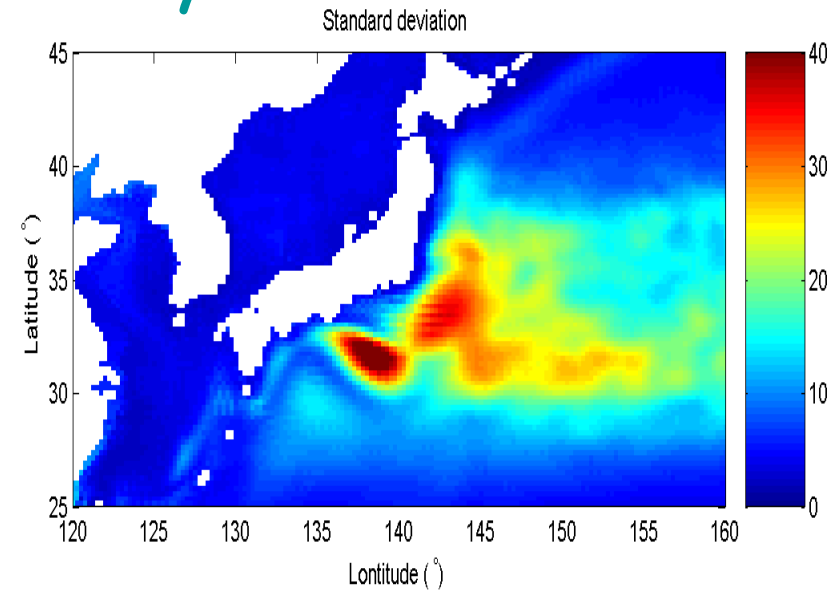
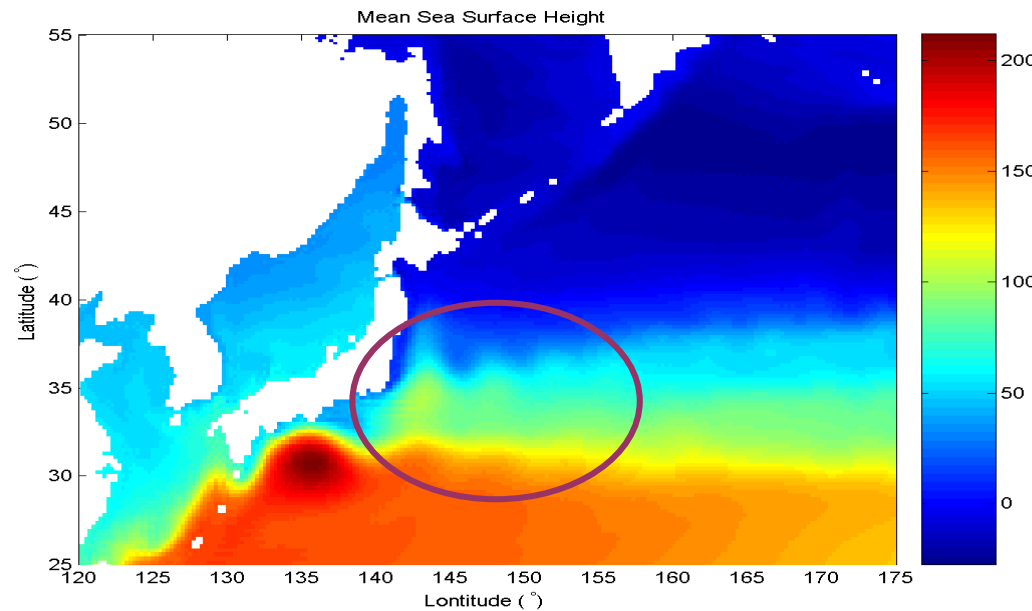


modelled ten-year standard deviation (year 41-50) of equivalent sea surface height (in cm)

Courtesy of Jim Richman (NRL)

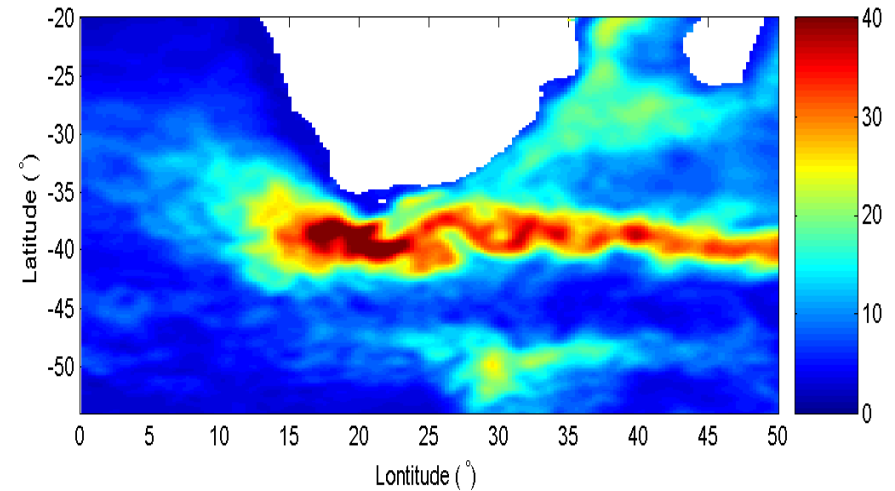
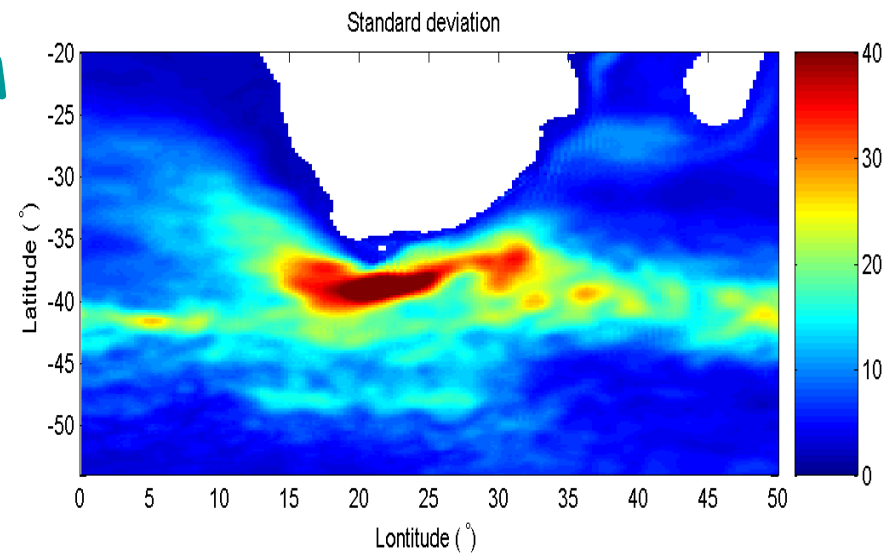
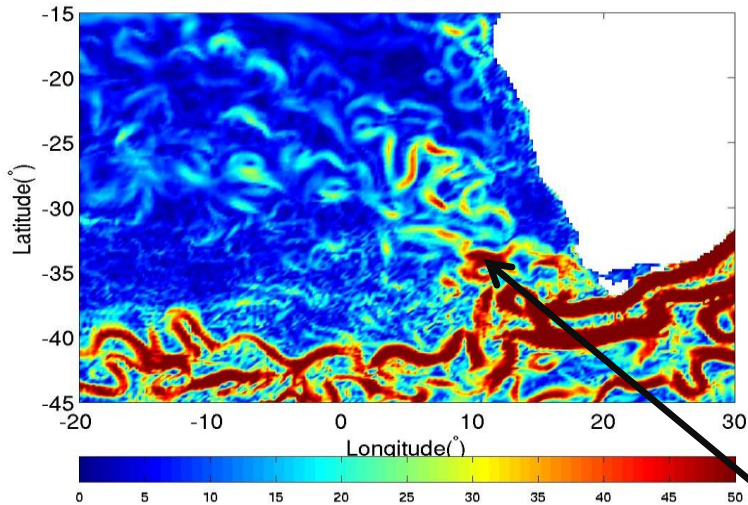
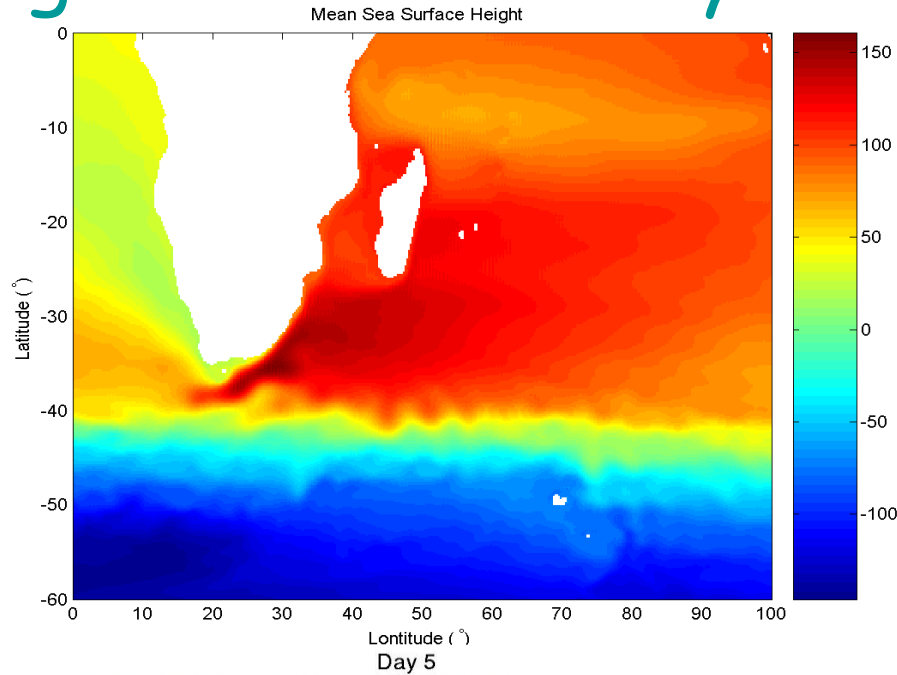


# Kuroshio Current System



Qiu and Chen (2005)

# Agulhas Current System



5-day averaged surface current speed (cm/s) in Cape Basin

Retroflection of the Agulhas Current

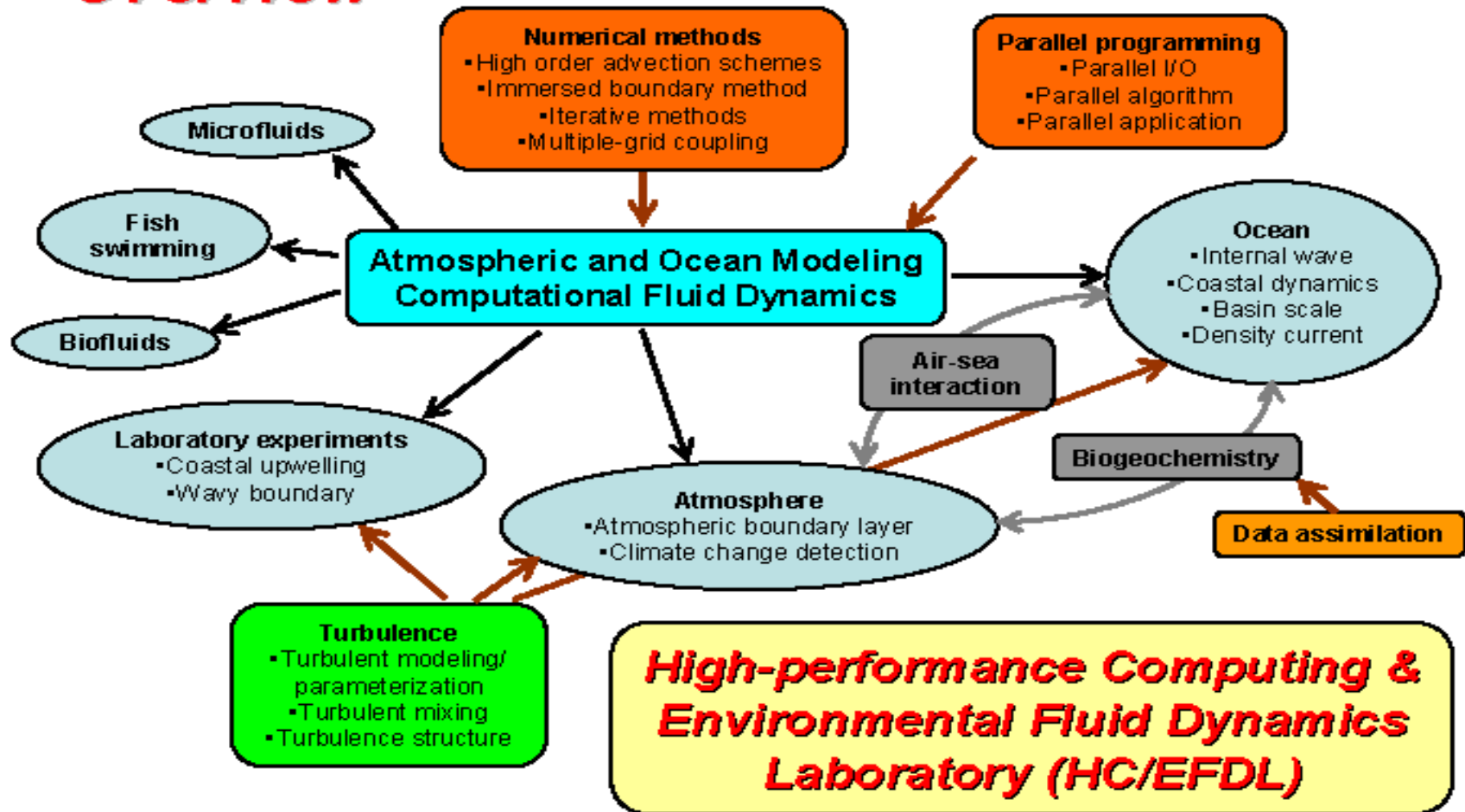


# Summary

- High-resolution Parallel Domain-decomposed TaIwan Multi-scale Community Ocean Model (PD-TIMCOM) is developed
  - Based on TIMCOM
  - Based on an efficient parallel EVP solver
  - Ideal (scalable) for parallel domain-decomposition
- Reasonable mean, standard deviation and skewness states
- Eddy-resolving global circulation patterns
- Fifty year simulation is almost completed
- Further validations and extremely high-resolution ( $1/16^\circ$ ) in Global Oceans and investigate the global ocean climate

# 高速計算與環境流體動力實驗室

## Overview



Questions.....?