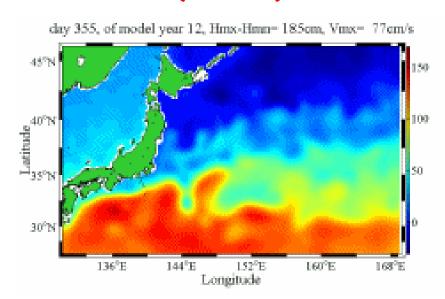
Development of a 1/16° Eddy-resolving Global Ocean Simulation within an Earth System Model Framework

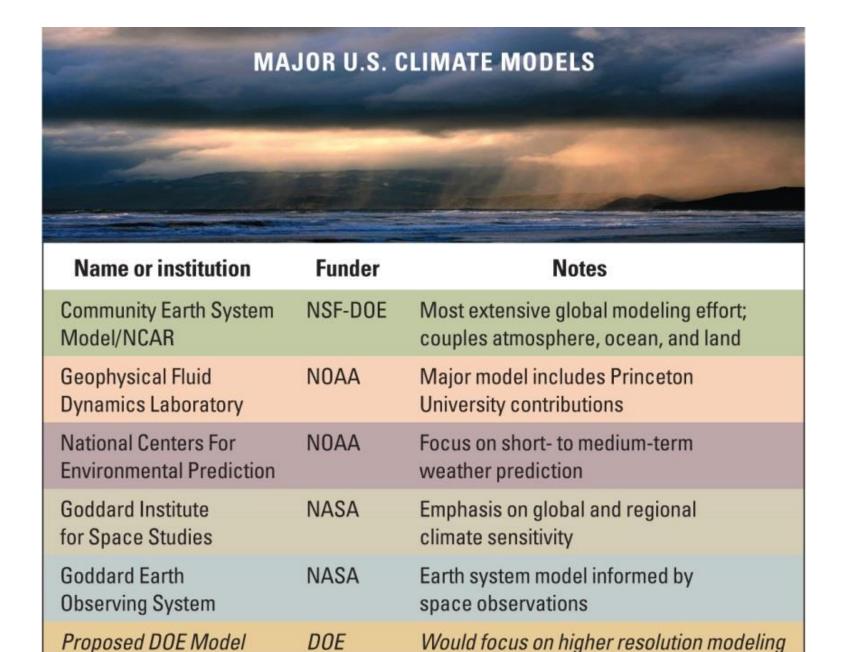
An international collaborative effort for Multi-scale Ocean Circulation System (MUSOC)



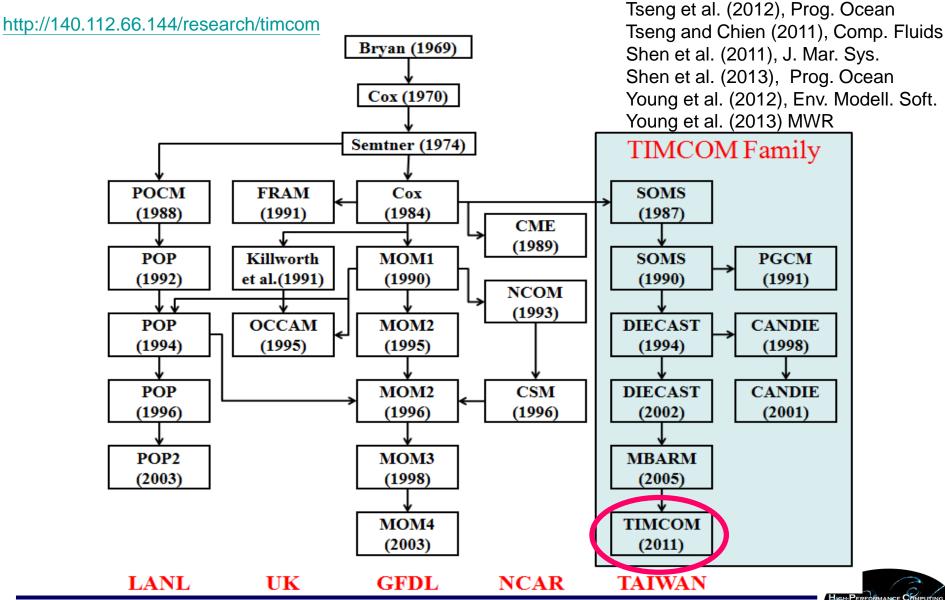
Yu-Heng Tseng, Luc Vandenbulcke, Chih-Chieh Young, Wen-ien Yu, Rui Caldeira, Mu-hua Chien, Yu-Chiao Liang

Objectives

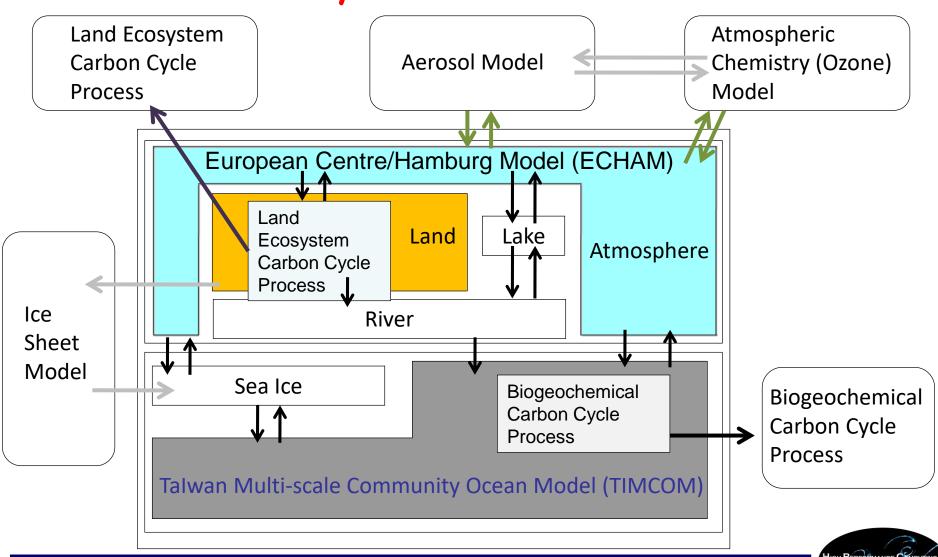
- Build and redesign a new ocean community model to resolve multi-scale dynamics with the most efficient twoway coupling in high accuracy
- Provide a scalable high-resolution global ocean circulation model (targeting at 1/60) for global ocean climate study in an Earth System Model Framework
- Operational and forecast capability while maintaining the simplest system structure
- International collaborative effort

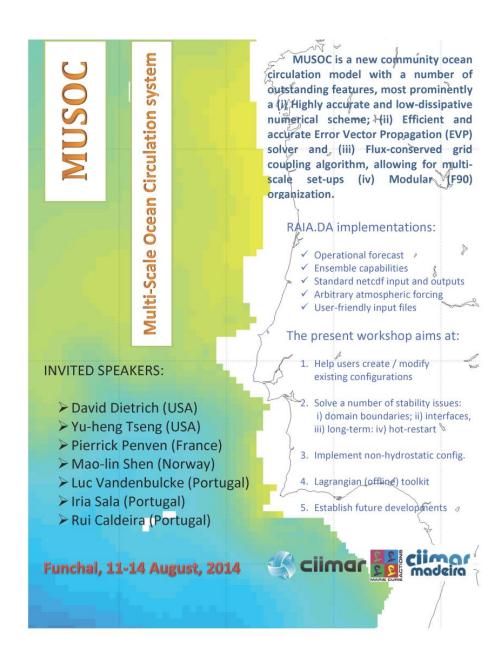


Multi-Scale Ocean Circulation System (PD-TIMCOM)



ECHAM/Taiwan Earth System Model





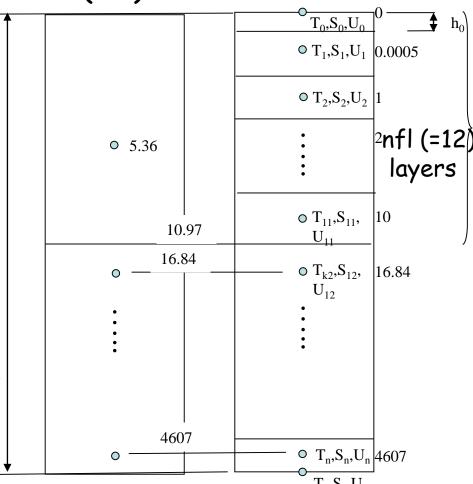
ECHAM/TWESM

(ECHAM5/SIT/TIMCOM)

TIMCOM (30)

SIT (41)



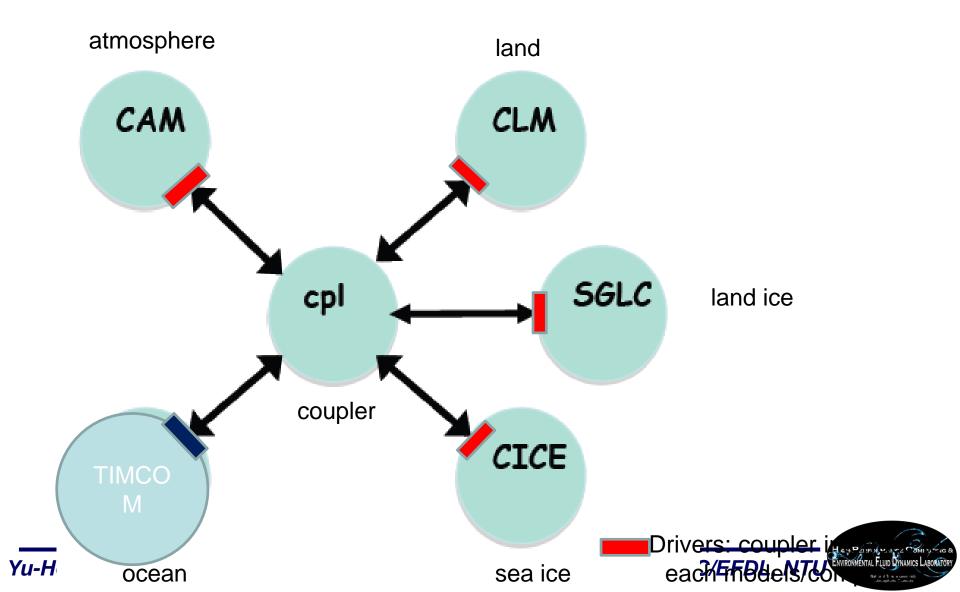


Improved MJO

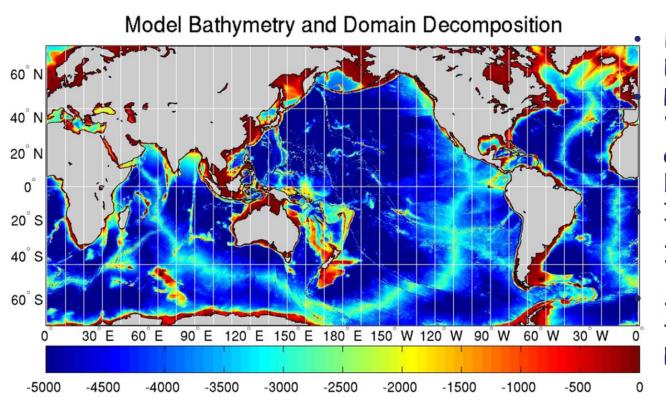
• Improved cool-skin simulation (Tu and Tsuang, 2005)

Seng

CESM framework--Drivers



Parallel Domain-Decomposed Taiwan Multi-Scale Community Model (PD-TIMCOM)



Bathymetry from 1-min Etopo1
Pacanowski and Philander Vertical mixing (1982) and Smagorinsky horizontal mixing (1993). Initial Temperature and Salinity from NOAA WOA 09.
Surface wind forcing from Hellerman and Rosenstein (1983)

- 1/16° and 1/4° horizontal resolution, latitudes covers from 72°S to 72°N. with 51 linear exponential levels vertically. (1440×792×51)
- Primitive, hydrostatic equation
- Fourth-order numerics combined Arakawa A and C-grid (1977)
- Rid-lid approx. and Free surface are used (Yang et al., in preparation)

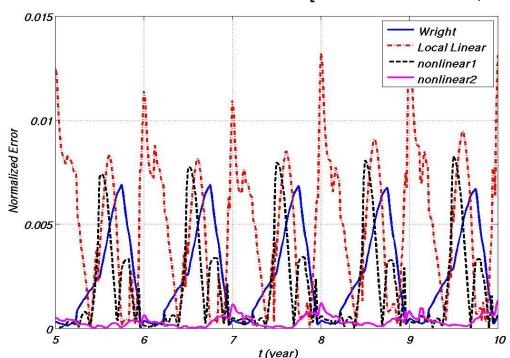
EQUATION OF STATE

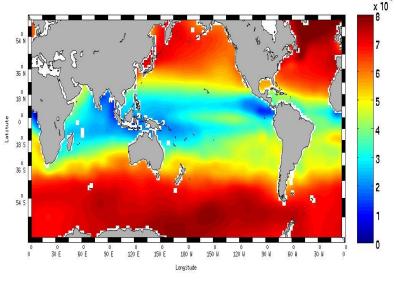
Wright's:
$$(\alpha - \alpha_0)(p + p_0) = \lambda \Rightarrow \rho = \frac{p + p_0}{\lambda + \alpha_0(p + p_0)}$$

Polynomial fit:

$$D(t, s, p) = C_0 + (C_t + C_{tt}t + C_{ts}s)t + C_s s$$

[Sanderson et al., 2001]





- UNESCO: standard solution.
- Error is normalized.
- Largest error of Wright's in autumn.
- Largest error of Local Linear in winter.
- · Largest error of nonlinear1 in summer.
- Largest error of nonlinear2 in winter.

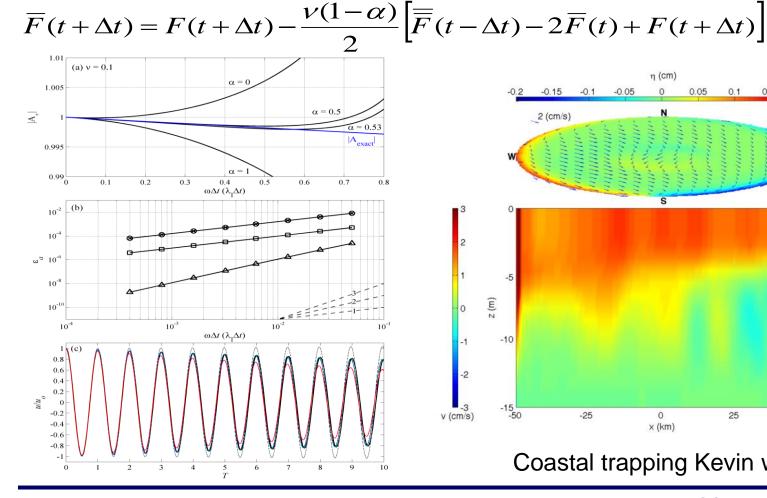
RAW-FILTER

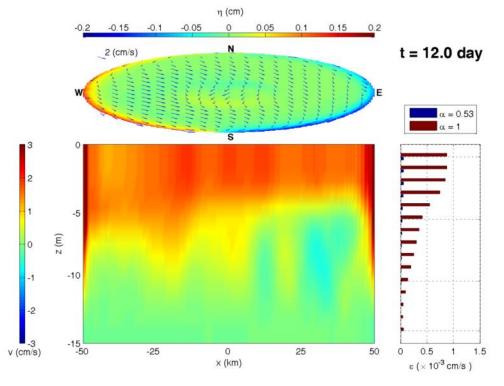
- 3rd-order accuracy in Amp. Young et al. (MWR in press)
- reduce time-splitting instability.
- easy to implement.
- · low computational cost.

$$\overline{\overline{F}}(t) = \overline{F}(t) + \frac{v\alpha}{2} \left[\overline{\overline{F}}(t - \Delta t) - 2\overline{F}(t) + F(t + \Delta t) \right]$$

•
$$\alpha = 0.53$$

$$\left[\begin{array}{l} \bullet \ \alpha = 1 -> RA\text{-}Filter \\ \text{[Williams, 2009]} \end{array}\right]$$





Coastal trapping Kevin wave

Numerical Experiments

- 1. 50 years TIMCOM 1/4° global simulation with WOA (2009) and HR (1984) wind forcing (rerun).
- 2. 50 years TIMCOM 1/4° global simulation with WOA09 and Scatterometer Climatology of Ocean Winds (SCOW, 2008) wind forcing (rerun).
- 3. 2004-2010 TIMCOM 1/4° global hindcast with NOGAPS daily forcing (in prep.).
- 4. 1960-2010 TIMCOM 1/4° global hindcast with ECMWF monthly forcing (in prep.).
- 5. 30 years TIMCOM 1/16° global simulation with Levitus (1994) and HR (1984) wind forcing (in proc.).
- 6. 30 years TIMCOM 1/16° degree global simulation with WOA09 and Scatterometer Climatology of Ocean Winds (SCOW, 2008) wind forcing (in prep.).
- 7. 2004-2010 TIMCOM 1/16° degree global hindcast with NOGAPS daily forcing (in prep.).
- 8. 1960-2010 TIMCOM 1/16° degree global hindcast with ECMWF monthly forcing (in prep.).

Current status

http://140.112.66.144/wiki/index.php/Timcom_Wiki

- . Focus mainly on the MEDINA and GLOBAL configurations
- For historical (climatological) studies of processes along the interface between Mediterranean Sea and Atlantic Ocean: promising (we see Meddies) but slow (not parallelized yet)
- For operational forecasts: being implemented now (not very much manpower to do it, so it's taking some time)

Doesn't really need parallel version for it (ensemble runs are parallel by themselves)



Recent additions

- Enhanced data input capacity (free-format ascii files, can read scalars, vectors, or load from netcdf files)
- Numbered models: input and output (000 to 999) for future ensemble runs
- Basic river runoff influences
- More flexible nudging configuration (incl. 3D coeffs.)
- Support for real time atmospheric forcing instead of climatological time: the model keeps track of the "modified julian day" (days after 18/november/1858).

Thus years have their exact number of days rather than 12*30=360 days



Recent additions

Surface fluxes for heat, momentum and freshwater

 Fluxes computed from atmospheric fields using bulk formulae (based on code from GETM model)
 or

Fluxes read from netcdf files

The module is initialized, data is read from netcdf files and interpolated in space during the model startup; and interpolated in time during model integration

Tested with data from NCEP, ECMWF and CORE-2

Data assimilation

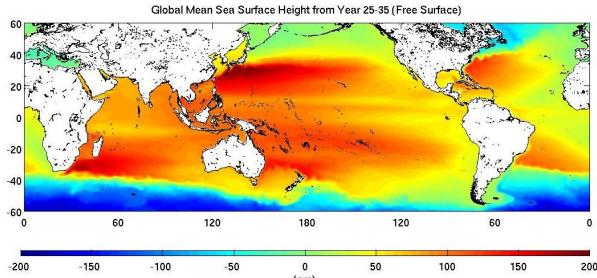
- Data assimilation independent from TimCOM
- Bash scripts alternatively run the model and the data assimilation code
- The data assimilation code modifies the TimCOM restart file
- . The data assimilation code we plan to use is OAK

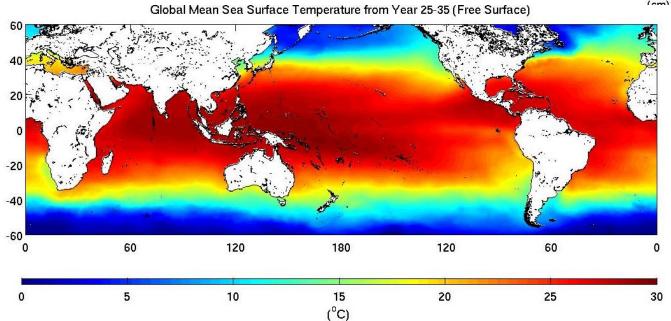
(modb.oce.ulg.ac.be/mediawiki/index.php/Ocean_Assimilation_Kit)

multi-grid, multi-variate, 3D, reduced-rank Kalman filter. The state-vector error covariance matrix can be obtained from a historical run ("fixed-base SEEK filter") or from an ensemble (EnKF)

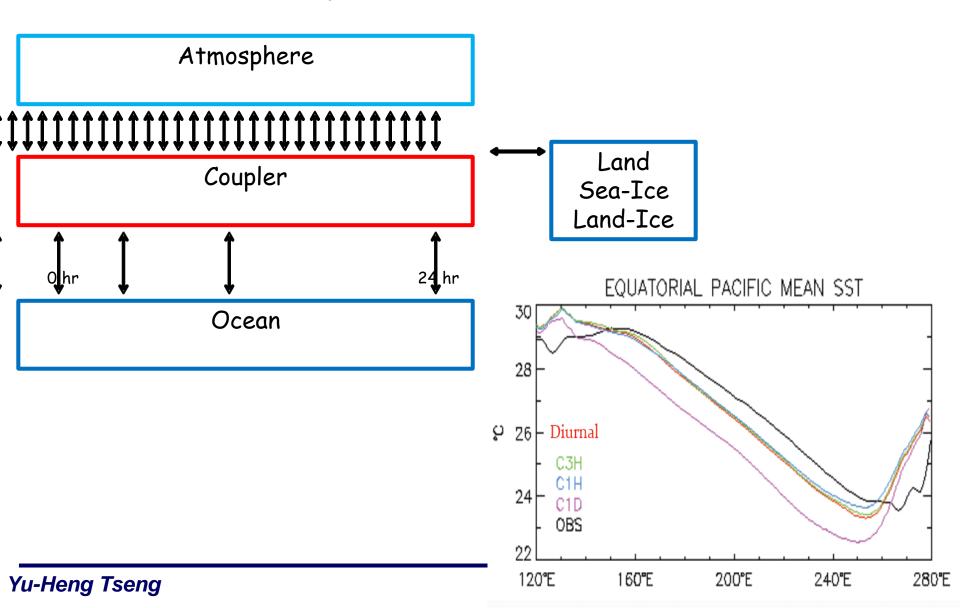


Mean SSH and SST

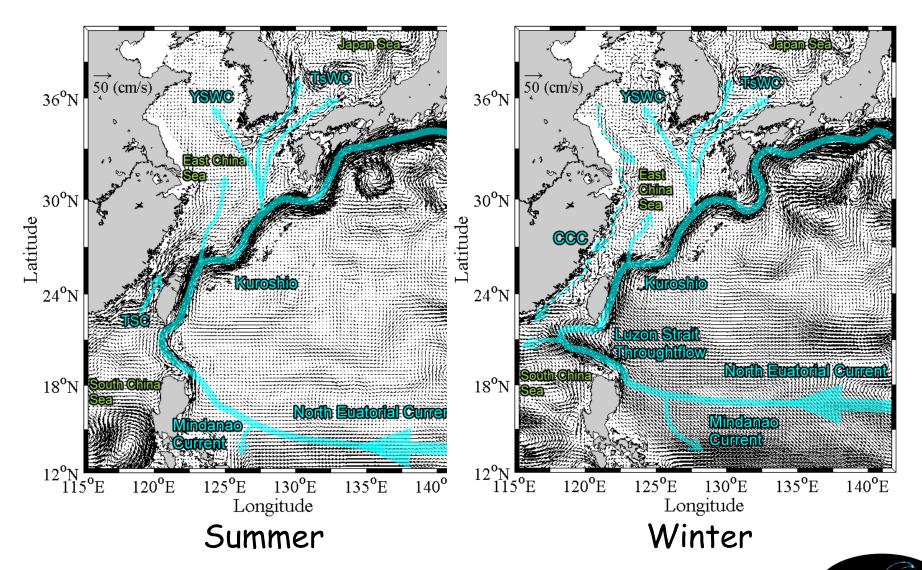




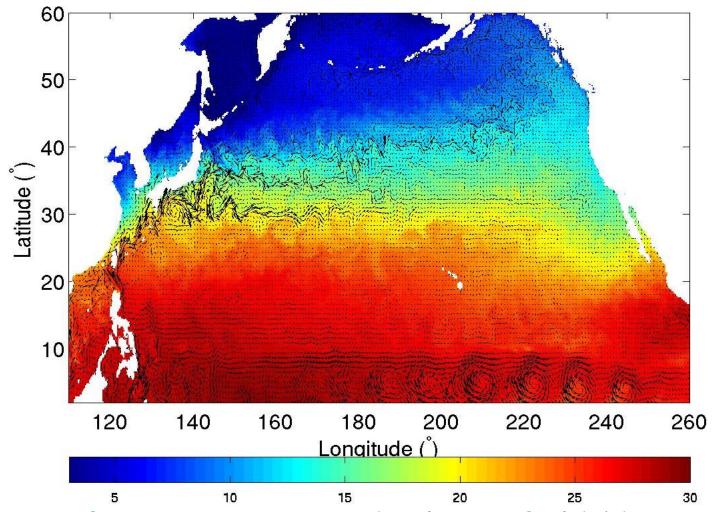
Future Plans



General seasonal circulation pattern-DUPOM



1/4° × 1/4° global resolution (domain 1442×720×26)



North Pacific temperature and velocity field (day 5, Year 49)

Yu-Heng Tseng

Governing Equations

λ: the longitudinal variable φ: the latitudinal variable z: the vertical variable

Continuity eqn.
$$\frac{1}{R\cos\varphi}\Big(\frac{\partial u}{\partial\lambda} + \frac{\partial(\nu\cos\varphi)}{\partial\varphi}\Big) + \frac{\partial\omega}{\partial z} = 0$$

Momentum eqn.

$$\begin{split} \frac{\partial u}{\partial t} &= -\mathcal{L}u + \Big(f + \frac{u\tan\varphi}{R}\Big)\nu - \frac{1}{\rho_0R\cos\varphi}\frac{\partial p}{\partial\lambda} + D_mu + \frac{\partial}{\partial z}\Big(A_u\frac{\partial u}{\partial z}\Big)\\ \frac{\partial v}{\partial t} &= -\mathcal{L}\nu - \Big(f + \frac{u\tan\varphi}{R}\Big)u - \frac{1}{\rho_0R}\frac{\partial p}{\partial\varphi} + D_m\nu + \frac{\partial}{\partial z}\Big(A_v\frac{\partial u}{\partial z}\Big) \end{split}$$

Conservation eqn. for temperature and salinity

$$\frac{\partial T}{\partial t} = - \mathcal{L}T + D_h T + \frac{\partial}{\partial z} \left(K_T \frac{\partial T}{\partial z} \right)$$

Egn. of State

$$\rho = \rho(S,T,p)$$

Hydrostatic Eqn. $\frac{\partial \mathbf{p}}{\partial \mathbf{p}} = -(\mathbf{p} - \overline{\mathbf{p}})\mathbf{g}$

$$D_{m(h)} = \frac{A_{m(h)}}{R^2} \left(\frac{1}{\cos^2 \varphi} \frac{\partial^2}{\partial \lambda^2} - \tan \varphi \frac{\partial}{\partial \varphi} + \frac{\partial^2}{\partial \varphi^2} \right) \qquad \mathcal{L} = \frac{u}{R \cos \varphi} \frac{\partial}{\partial \lambda} + \frac{v}{R} \frac{\partial}{\partial \varphi} + \omega \frac{\partial}{\partial z}$$